





SUBJECT AREAS:

MINERALOGY

SOLID EARTH SCIENCES

STRUCTURAL GEOLOGY

SURFACES, INTERFACES AND
THIN FILMS

Received 25 November 2012 Accepted 15 March 2013 Published

2 April 2013

Correspondence and requests for materials should be addressed to A.Z. (azaoui@polytech-lille.fr)

# Nanoscale analysis of the morphology and surface stability of calcium carbonate polymorphs

W. Sekkal & A. Zaoui

Civil Engineering and geo-Environment Laboratory LGCgE- (EA 4515) Lille Nord de France, Polytech'Lille. Université Lille 1 Sciences et Technologie. Cité Scientifique. Avenue Paul Langevin. 59655 Villeneuve D'Ascq Cedex, France.

Under earth surface conditions, in ocean and natural water, calcium carbonate is ubiquitous, forming anhydrous and hydrous minerals. These hydrous phases are of considerable interest for their role as precursors to stable carbonate minerals. Atomistic simulation techniques have been employed here to perform a comprehensive and quantitative study of the structural and energetic stability of dry and hydrous surfaces of calcium carbonate polymorphs using two recently developed forcefields. Results show that the dry forms are prone to ductility; while hydrous phases are found to be brittle. The (001) surface of monohydrocalcite appears to be the most stable (0.99 J/m²) whereas for the ikaite phase, the (001) surface is the most stable. The corresponding value is 0.2 J/m², i.e. even lower than the surface energy of the (1014) surface of calcite. Beautiful computed morphology pictures are obtained with Xiao's model and are very similar to the observed SEM images.

arbonates are among the most abundant minerals in the environment. We can found them in aquatic systems forming huge amounts of scale and ocean sediment1, in the atmosphere2 and as part of the sediment and rock record<sup>3</sup>. These minerals affect the chemistry of aquatic systems by regulating pH and alkalinity through dissolution/precipitation equilibrium. Moreover, they can govern the mobility and cycling of metal contaminants and radionuclides via ion exchange, adsorption, and co-precipitation reactions<sup>4</sup>. Calcium carbonate (CaCO<sub>3</sub>) is one of the constituents of argillite used for the wastes disposal. It has an important role in restraining the swelling properties of smectitic layers and produce hydraulic and mechanical reinforcement of the clay phase. It is one of the most abundant materials found in sedimentary rock in all parts of the Earth surface (it makes up 4% of the earth's crust). CaCO<sub>3</sub> occurs naturally in six different forms: three crystalline polymorphs, calcite, aragonite, and the metastable vaterite; two hydrate phases (monohydrocalcite (CaCO<sub>3</sub>-H<sub>2</sub>O) and Ikaite (CaCO<sub>3</sub>-6H<sub>2</sub>O)); and amorphous calcium carbonate (ACC). The monohydrocalcite structure has 8-fold coordinated calcium ions. Water is situated between the screw axes (parallel to the c-axis), bound to the carbonate ions via hydrogen bridges<sup>5</sup>. Additionally, monohydrocalcite is found to be metastable under all conditions<sup>6</sup>. However, it does form in some predominantly freshwater environments7. The structure of ikaite is characterized by hydrated calcium carbonate chains, aligned along the a- axis and are linked together by hydrogen bonds; the calcium ions are 8-fold coordinated8. It is found that ikaite is stable at high pressure and low temperatures in an alkaline aqueous environment<sup>9-11</sup>. It can be found in bottom waters in fjords and on deep sea floors<sup>12</sup>.

Calcite is one of the most stable phase at atmospheric pressure and temperature, while aragonite is found to be the most stable polymorph in high-pressure environments, although it is possible for aragonite to form at atmospheric pressure in certain conditions, such as hot springs<sup>13</sup>. Vaterite, the metastable form of calcium carbonate, changes to form another polymorph, or is stabilized by water adsorption on its surfaces<sup>14</sup>. We can explain the differences between the three polymorphs from the different arrangement of the CO<sub>3</sub> ions in their crystal structures.

Many forcefields have been developed to describe the CaCO<sub>3</sub> interactions<sup>14–22</sup>. Other empirical models have been used to explore the bulk and surface properties of CaCO<sub>3</sub> at ambient conditions<sup>18,23–25</sup>. Parker et al.<sup>26</sup>, using an atomistic model, examined the surface precipitation and dissolution processes. On the other hand, Fisler et al.<sup>22</sup> used the shell model to simulate the rhombohedral carbonates and their point defects. Archer et al.<sup>17</sup> developed a new empirical potential to study the first-order phase transition between calcite I and calcite II. Recently, new models have been developed describing the interaction of the system with water, and even organics, with a view to addressing biomineralization. Raiteri et al.<sup>20</sup> proposed a new potential model, which gives successful



atomic-resolution simulations of the growth of amorphous calcium carbonate. However, the forcefield reproduces well structural and thermodynamic properties of CaCO<sub>3</sub> but not the elastic behavior. The latest forcefield published by Xiao et al.<sup>21</sup> for CaCO<sub>3</sub> derived a new set of forcefield parameters, which reproduces accurately the elastic and shear modulus for bulk aragonite crystal.

#### Results

The broad objective in this paper is to use, for the first time, two latest forcefields developed recently by Raiteri $^{20}$  and Xiao $^{21}$  in order to find out the best model to analyze the surface of dry and hydrous calcite and its polymorphs. X-Rays diffraction patterns will be evaluated and compared to experimental findings. Structural and mechanical quantities will be evaluated for calcite, aragonite, vaterite, monohydrocalcite and Ikaite phases. We will discuss about the structural and energetic stability of CaCO $_3$  polymorphs through the calculation of the surface and attachment energies as well as the equilibrium morphologies.

**Bulk properties of calcium carbonate polymorphs.** Calculations are performed using the General Utility Lattice Program 4.0 (GULP)<sup>27</sup>. Internal energy minimizations have been done in NPT ensemble allowing all individual ionic coordinates and lattice parameters to vary. The unit cell was fully optimized in order to obtain the equilibrium crystal structure. The search of the local minima adopts the Newton-Raphson procedure, with the Broyden-Fletcher-Gold-farb-Shanno (BFGS) scheme to update the Hessian. Structural

parameters of the known anhydrous (calcite, aragonite, vaterite) as well as hydrous polymorphs of CaCO<sub>3</sub> (monohydrocalcite and Ikaite), have been evaluated with both models (Raiteri's and Xiao's). They are compared to experimental data taken from the American Mineralogist database. The absolute errors of the lattice parameters a, b, c, and equilibrium volume are displayed in Figure 1 (a, b, c, d). We can see that for the case of calcite, the Raiteri's model reproduces better the structure with a lower error (less than 2%). However, when studying aragonite or vaterite structures and the hydrous phases (monohydrocalcite and Ikaite), our results show that the Xiao's forcefield reproduces well the structure of hydrated CaCO3 minerals, since the deviations of the unit cell parameters are less than 3%. Figure 1 (d) illustrates that the volume evaluated with Xiao model is obtained with an error less than 4.5% for the dry and hydrated phases, while Raiteri model gives a higher error (7-8%) for monohydrocalcite and ikaite and around 4% for the dry phases.

In Figure 2, we plot our calculated X-Ray diffraction patterns ( $\lambda=1.54\text{Å}$  for calcite carbonate polymorphs using Raiteri's and Xiao's forcefield, compared with experimental findings. As one would expect, the Raiteri's forcefield reproduces perfectly the calcite phase than Xiao's model (Figure 2 (a)). This confirms our structural data as discussed above. We can see that from  $2\theta=70^\circ$ , the curve obtained from Xiao forcefield does not match the experimental findings. For the case of aragonite phase (Figure 2 (b)), the obtained X-Rays patterns from Xiao model are very similar to the experimental curve. We note a shift of the Raiteri's curve along x, which is more pronounced from  $2\theta=30^\circ$  to  $42^\circ$ . For the case of vaterite, the

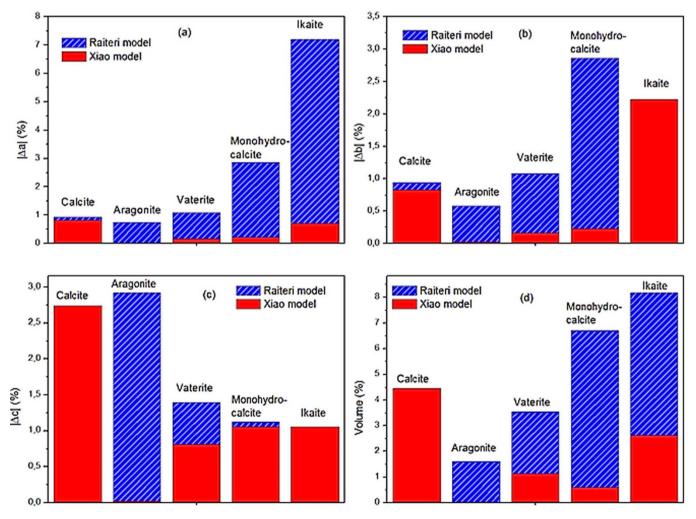


Figure 1 The absolute errors of the lattice parameters and volumes for CaCO<sub>3</sub> phases.



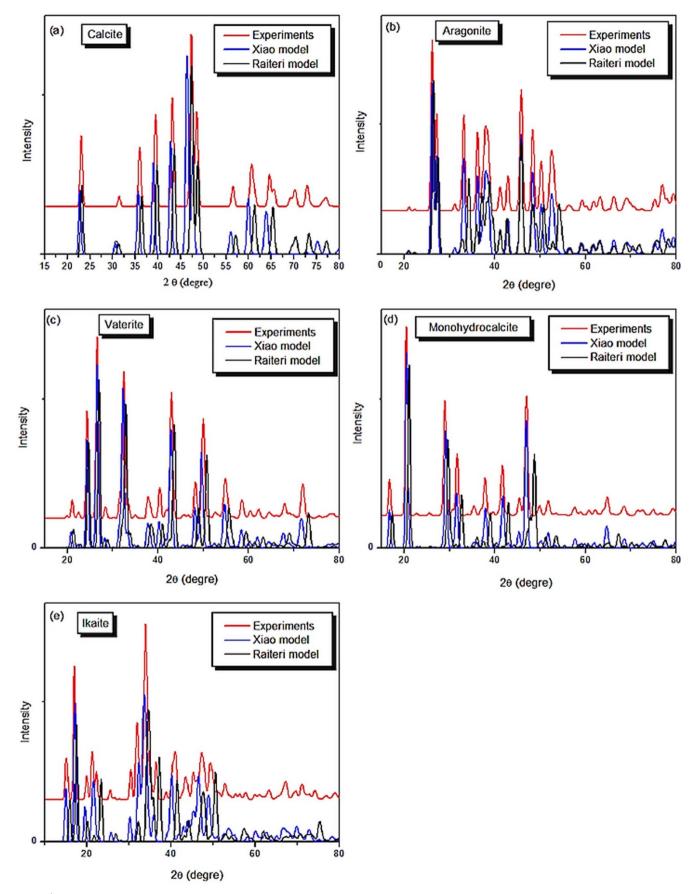


Figure 2 X-Rays diffraction patterns for dry and hydrous CaCO<sub>3</sub> polymorphs calculated with Raiteri's and Xiao's forcefields and compared to experimental findings: (a) calcite, (b) aragonite, (c) vaterite, (d) monohydrocalcite, (e) Ikaite.



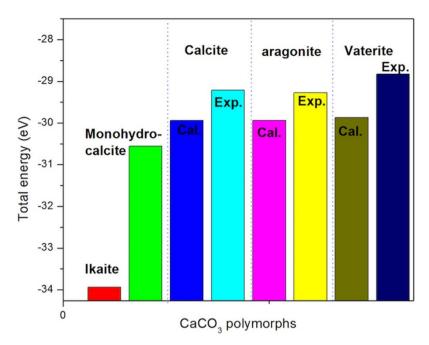


Figure 3 | Energetic stability of CaCO<sub>3</sub> polymorphs.

monohydrocalcite and ikaite phases, X-Rays diffraction patterns obtained from Xiao's model are similar to experimental curves (Figure 2 (c), (d), (e)). In Figure 3, the energetic stability of  ${\rm CaCO_3}$  polymorphs is presented compared to experimental measurements²8. We can see that ikaite has the lowest total energy (-33.9393 eV). Calcite (-29.9371 eV) and aragonite (-29.9358 eV) phases have a small energy difference, about 0.0013 eV. Vaterite with a total energy equal to -29.8682 eV is the less stable phase.

Concerning the mechanical properties, the elastic constants tensor is calculated from the second derivative matrix at the optimized structures and is listed in Table 1(a) for the hydrated phases. In a recent work of Xiao et al. the elastic data, for the dry phases, were reported and compared with experiments  $^{21,29,30}$ . We can notice that  $C_{11}$  for Ikaite is smaller than  $C_{22}$  and  $C_{33}$  what means that this structure is more compressive in the direction perpendicular to the layers (layers parallel to (100)). We have also calculated the isotropic bulk (B) and shear (G) moduli obtained from the Voigt-Reuss-Hill averages  $^{31}$ . We can define B as the amount of pressure required for a unit volume change of the material; G measures the shear forces required to induce a unit shape change measured, while E is the force per unit of cross-sectional area needed for a unit dimensional change.

Table 1a   Elastic data for monohydrocalcite and Ikaite						
	Monohyo	Monohydro calcite		iite		
	а	b	а	b		
C <sub>11</sub> (GPa)	99.3	98.8	41.7	59.2		
C <sub>22</sub> (GPa)			96.1	75.4		
C <sub>33</sub> (GPa)	91.9	88.1	82.0	62.7		
C <sub>12</sub> (GPa)	25.6	28.9	18.8	20.4		
C <sub>13</sub> (GPa)	34.8	42.2	27.8	26.3		
C <sub>23</sub> (GPa)			18. <i>7</i>	16.8		
C <sub>44</sub> (GPa)	52.8	32.1	8.0	4.2		
C <sub>55</sub> (GPa)			8.9	2.7		
C <sub>66</sub> (GPa)	36.8	34.9	16.2	12.9		

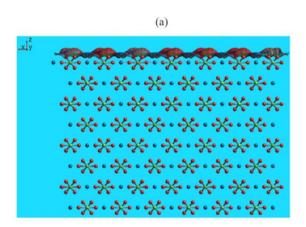
In Table 1(b) , we have listed the bulk modulus, B, shear modulus, G, Young's modulus, E and B/G for all phases. Using the ratio B/G (the resistance to fracture (B) relative to the plastic deformation resistance (G)) introduced by Pugh $^{32}$ , we can analyze the ductility of a material. A value of B/G higher than 1.75 indicates a tendency for ductility, while a value of B/G lower than 1.75 indicates a tendency for brittleness. The dry forms of calcium carbonate yield values higher than 1.75, suggesting that calcium carbonate materials are prone to ductility. Moreover, the value of B/G on going from CaCO $_3$  to CaCO $_3$ -H $_2$ O and CaCO $_3$ -6H $_2$ O decreases, which suggests that the inclusion of water in the crystal makes the material more brittle. In fact, the bulk modulus decreases with about 36% from calcite to monohydrocalcite and about 56–82% from calcite to ikaite.

**Surface analysis.** Starting from the experimental structure (a = b = 4.988 Å, c = 17.061 Å;  $\alpha = \beta = 90^{\circ}$  and  $\gamma = 120^{\circ}$  the optimized bulk structure of calcite was cut to obtain the surfaces  $(10\overline{14})$ ,  $(11\overline{20})$ ,  $(10\overline{10})$ ,  $(10\overline{11})$  Ca<sup>+2</sup>,  $(10\overline{11})$  CO<sub>3</sub><sup>-2</sup>,  $(01\overline{12})$  (see Table 2). Cleavage along the  $\{10\overline{14}\}$  plane gives rise to the characteristic rhombohedra shape of calcite crystals. This plane contains both Ca<sup>+2</sup> and CO<sub>3</sub><sup>-2</sup> ions, making the surface charge neutral (Figure 4 (a)), and thus stable with a low surface energy. Furthermore, it has also a higher density of ions compared to other possible neutral planes. For instance, the  $(01\overline{12})$  and the  $(10\overline{11})$  planes are polar as they are terminated by either a layer of calcium or carbonate ions, leading to a positively or negatively charged surface respectively (Figure 4 (b), (c)). We plot in the top of the slabs the iso-surfaces. These are obtained using the Gaussian description of molecular shape. It is similar to the idea of rolling a probe atom over the surface of the system being studied.

For the aragonite phase, the crystal was cut to obtain the following low index surfaces: (010), (100), (001), (110), (011), (101), and (111). The (010) and (110) surfaces are cleavage planes while the (110) is also a twinning plane (see Table 3). The surfaces can be cut in more than one way, giving rise to planes terminated by either calcium ions or oxygen atoms of the carbonate groups. In the case of the metastable vaterite phase, all surfaces have both calcium and carbonate terminated planes as presented in Table 4. Surfaces are allowed to relax, the calculated surface and attachment energies with both



		Calcite			Aragonite		Vate	erite	Monohyd	ro calcite	Ikc	aite
	а	b	С	а	b	d	а	Ь	а	b	а	b
В	83.1	85.7	80.6	72.6	67.7	67.1	88.1	69.1	53.4	44.4	36.3	15.6
G	32.6	35 <i>.</i> 7	32.7	38.8	35.8	35.8	36.9	18.8	36.9	27.6	31.1	13. <i>7</i>
3/G	2.5	2.4	2.4	1.8	1.9	1.8	2.3	3.6	1.4	1.6	1.16	1.13
-×	108.6	78.1		185. <i>7</i>	132.7		-75.0	58.3	80.4	<i>7</i> 5.1	38.4	45.4
- y	108.6	78.1		81.0	69.3		-75.0	58.33	80.4	<i>7</i> 5.1	96.4	66.8
=' -z	54.62	39.45		58.56	76.95		145.9	110.6	72.52	60.2	81.9	44.2



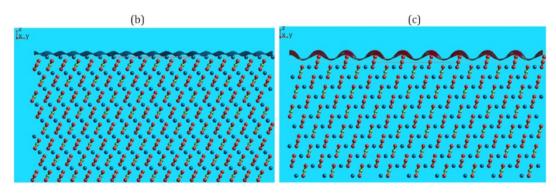


Figure 4 | (a)  $(10\overline{1}4)$  surface of calcite; on the top are presented iso-surfaces. (b)  $(01\overline{1}2)$  surface of calcite with Ca termination, (c)  $(10\overline{1}1)$  surface of calcite with CO<sub>3</sub> termination. Blue atoms represent calcium, green atoms are carbon and red atoms represent oxygen.

	Sur	face energy (J.m <sup>-2</sup> )		Attachment energy (eV)		
Surface (Hexagonal indices)	Raiteri's forcefield	Xiao's forcefield	Other results <sup>a</sup>	Raiteri's forcefield	Xiao's forcefield	Other results <sup>a</sup>
$\overline{(10\overline{1}4)}$	0.71	0.51	0.59	-11.34	-8.69	-7.82
(1120)	1.11	0.76	1.39	-33.91	-24.58	-30.24
$(10\overline{10})$	1.04	0.64	0.97	-33.29	-14.37	-78.86
(10 <u>1</u> 1)Ca	1.61	1.63	1.23	-11.92	-30.63	-31.89
(1011)CO <sub>3</sub>	1.36	0.97	1.14	-17.88	-24.58	-28.72
$(01\overline{12})$	4.85	1.56		-206.70	-100.78	



		Surface energy (J.m <sup>-2</sup> )	)	Attac	chment energy (eV)	
Surface	Raiteri's forcefield	Xiao's forcefield	Other results <sup>a</sup>	Raiteri's forcefield	Xiao's forcefield	Other results
(010) Ca	1.14	0.80	0.96	-4.62	-3.20	-10.91
(010) CO <sub>3</sub>	1.57	1.10	1.50	-11.96	-10.48	-32.68
(100)	1.71	1.23	1.50	-15.95	-10.75	-39.58
(001) Ca	1.59	1.36	1.05	-12.73	-9.58	-35.41
001) CO <sub>3</sub>	1.01	0.78	0.85	-5.35	-3.92	-12.46
(110) Ca	1.03	0.73	0.88	-8.61	-6.10	-10.50
110) CO <sub>3</sub>	1.30	0.92	1.04	-28.14	4.37	-31.45
101) Ca	1.19	0.90	0.99	-10.54	-6.18	-15.11
101) CO <sub>3</sub>	1.31	1.00	1.08	-14.52	-12.03	-16.14
011) CO <sub>3</sub>	0.90	0.84	0.69	-5.97	-9.08	-6.38
$011) CO_3$	1.09	1.33	1.16	-14.33	-16.37	-17.40
011) CO <sub>3</sub>	1.33	1.20	0.99	-10.08	-16.23	-10.61
111) Ca	1.30	0.94	1.40	-23.39	-3.82	-33.19
111) CO <sub>3</sub>	1.09	0.83	0.84	-14.94	-6.50	-18.38

	Surface energy (J.m <sup>-2</sup> )			Attachment energy (eV)			
Surface	Raiteri's forcefield	Xiao's forcefield	Other studies <sup>a</sup>	Raiteri's forcefield	Xiao's forcefield	Other studies a	
(010) Ca	1.51	1.34	1.35	-92.65	4.04	-35.58	
(010) CO <sub>3</sub>	0.75	0.75	0.62	-37.48	-37.48	-8.29	
(100) Ca	1.55	1.34	1.31	-115.51	4.04	-17.50	
(100) CO <sub>3</sub>	0.75	1.54	1.39	-37.48	-91.41	-14.23	
(110) Ca	1.54	0.87	1.18	-146.66	-107.35	-25.94	
(110) CO₃	0.87	0.98	1.04	-49.33	-55.29	-15. <b>47</b>	
(011) Ca	1.29	1.16	1.22	-41.10	-32.70	-59.90	
(011) CO <sub>3</sub>	1.12	1.12	0.93	-55.23	-55.23	-19.37	
(101) Ca	1.35	0.86	1.47	-34.66	-20.14	-19.33	
(101) CO <sub>3</sub>	1.22	0.79	1.07	-43.46	-16.62	-18.29	
(111) Ca	1.48	0.78	0.82	-190.73	-109.70	-19.47	
(111) CO₃	2.39	0.86	0.85	-125.23	-100.57	-25.47	

forcefields are given in Tables 2, 3, and 4 and are compared to other theoretical results<sup>14</sup>.

#### **Discussion**

In this part we will discuss about the best forcefield for modeling the surface of calcite and its polymorphs. We will analyze, for the first time the equilibrium morphology of dry and hydrous phases of CaCO<sub>3</sub>.

In the case of calcite (Table 2), it is clear that the relaxed  $(10\overline{1}4)$  surface is the most stable surface, with a lowest surface value and attachment energy. The corresponding values are found to be equal to 0.51 J/m² and -8.69 eV. We notice also that Xiao's forcefiled gives lower values, which agree better with the experiment and other

Table 5 | Relaxed surface and attachment energy of monohydro-

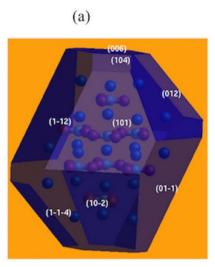
caiciic	•			
	Surface ene	Attachment e	energy (eV)	
(hkl)	Raiteri's model	Xiao's model	Raiteri's model	Xiao's model
(100)	1.48	1.54	-13. <i>7</i> 5	-7.82
(010)	1.39	1.21	-10.94	-6.62
(001)	1.00	0.99	-24.94	-5.66

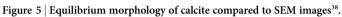
theoretical data<sup>33–35</sup>. The other polar surfaces are unstable by themselves and will either reconstruct to neutralize the dipole moment or adsorb ions to remove the dipole moment. However, they can be stabilized when grown on charged substrates.

Concerning the surface of the aragonite phase and according to the above structural calculations, we remark (Table 3) that the surface energies obtained with Xiao's model agree better with the other theoretical works<sup>14</sup>. In the other hand, we can see that the surface energy of the pure (010) surface quoted in Table 3 shows the calcium terminated plane to have a lower surface energy than the carbonate surface and hence to be the most stable surface plane. This is due to the rotation of carbonate groups rotate, which tend to lie flat in the surface rather than upright as in a bulk termination. On the other hand the calcium-terminated plane shows hardly any relaxation. In

Table	Table 6   Relaxed surface and attachment energy of Ikaite							
	Surface ene	Attachment energy (eV)						
(hkl)	Raiteri's model	Xiao's model	Raiteri's model	Xiao's model				
(100) (010) (001)	1.46 0.35 0.23	0.52 0.37 0.20	-13.65 -7.48 -3.04	-8.19 -5.79 -2.57				

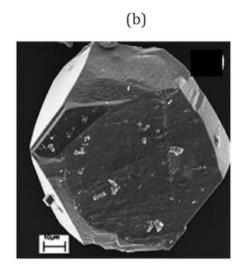






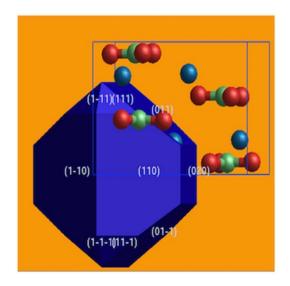
addition, the attachment energies agree well with the surface energies, where we notice that the calcium plane is favored over the carbonate plane through its low value indicating a small growth rate. The same remarks for the (110) surface. Concerning the (011) surface, it has different cuts with one or more surface oxygen ions. However, their surface energies vary greatly as shown in Table 3 from a very stable surface with  $\gamma=0.84~J~m^{-2}$  to the least stable with a surface energy of 1.33 J m $^{-2}$ . We notice that the attachment energies show the same trend as the relaxed surface energies with a very low value for the most stable plane. Consequently, this surface would be expressed in the growth morphology.

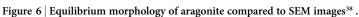
For the relaxed surface of vaterite phase, we note that our results based on Xiao's model are generally in good agreement with the theoretical calculations<sup>14</sup>. It is clear from Table 4 that the (010) carbonate plane is the dominant surface, with the lowest surface and attachment energies of the series. Moreover, the (100), (001), and (110) surfaces have relatively large surface and attachment energies, even if some surfaces have been relaxed sufficiently. We remark that the (011) calcium plane and both (101) surfaces are dipolar and both need extensive relaxation on energy minimization.

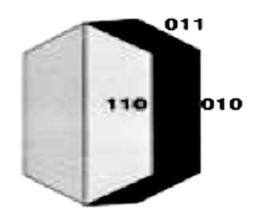


In Table 5 and 6, we present our results for the relaxed surface of hydrated CaCO<sub>3</sub>, monohydrocalcite and Ikaite, using Raiteri's and Xiao's forcefields. For Monohydrocalcite (CaCO<sub>3</sub>-H<sub>2</sub>O), the surface energies found using Xiao' model are 1.54, 1.21 and 0.99 J/ m<sup>2</sup> for the x, y, and z surfaces, respectively (Table 5). The corresponding attachment energies are -7.82, -6.62 and -5.66 eV for the x, y, and z surfaces, respectively. These results show that the (100) surface is less stable than the other two surfaces, which implies that the (100) surface will not appear in the monohydrocalcite equilibrium morphology. For the case of relaxed Ikaite surface, the obtained surface energies (Table 6) are equal to 0.52, 0.37, 0.20 J/m<sup>2</sup> for (100), (010) and (001) planes, respectively. The corresponding attachment energies are -8.19, -5.79, -2.57 eV for (100), (010) and (001) planes, respectively. We can clearly see that the (001) surface is the most stable (0.2 J/m<sup>2</sup>), with a value lower than that found for the most stable  $(10\overline{1}4)$ surface of calcite However, the (100) plane with 0.52 J/m<sup>2</sup> would not be expressed in the growth morphology.

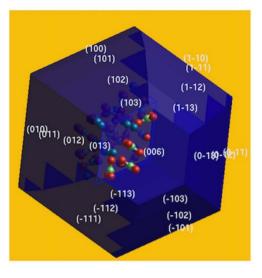
In order to measure the relative stability of the surfaces, it is necessary to compute the equilibrium morphology of a crystal, which is



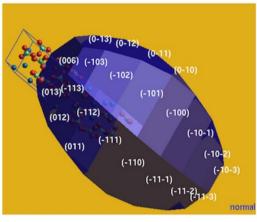












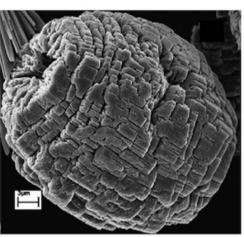


Figure 7 | Equilibrium morphology of vaterite compared to SEM images  $^{38}$  .

determined by the surface energy and the related growth rate for the various surfaces. The Gibbs-Wulff theory provides a simple mathematical approach that predicts the equilibrium crystal shape, which should possess a minimal total surface energy for a given volume<sup>36,37</sup>. This induces that a surface with a high surface free energy has a large growth rate, and this fast growing surface will not be expressed in the equilibrium morphology of the resulting crystal. Therefore, we consider only surfaces with low surface energies and hence slow growth will be expressed. It is well known that at 0 K, the surface free energy is a close approximation of the surface energy as calculated by static lattice simulations. In fact, this is due to the entropy term, which is found small when included in the surface free energy. Thus, the surface energies can be assumed to determine the equilibrium morphology of the crystal.

Figures 5, 6 and 7 show the equilibrium morphologies of dry calcium carbonate (calcite, aragonite, vaterite) obtained for all allowed faces and compared to the experimental images of scanning electron microscopy (SEM)<sup>38</sup>. We remark that our computed morphologies are very similar to the experimental ones. For the hydrated phases, we show in Figure 8 (a) the equilibrium morphology of Ikaite compared to the SEM image<sup>39</sup>. In Figure 8(b) the predicted equilibrium morphology for monohydrocalcite phase is presented since no experimental images are available. These results confirm once again the potential models employed and support strongly our findings.

#### **Methods**

**Potential model and surface simulation.** In Raiteri's model<sup>19</sup>, the oxygen-oxygen and calcium-oxygen interactions are modeled via a Buckingham potential. The covalency within the  $\mathrm{CO}_3^{2-}$  group is modeled by a Morse potential plus a three-body potential and a torsional potential function to restrain the planar  $\mathrm{CO}_3^{2-}$  group from folding. The interaction between the calcium ions and water is represented by the Lennard-Jones (12,6) potential including molecular flexibility<sup>20</sup> using the SPC/Fw model<sup>10</sup>.

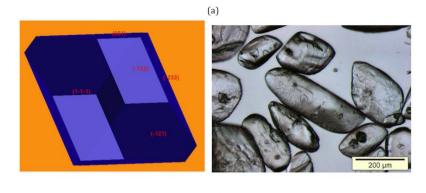
Xiao et al.<sup>21</sup> proposed a new forcefield to describe interatomic interactions for the bulk aragonite crystal in an aqueous and/or protein environment using the TIP3P water model<sup>41</sup>. In comparison to Raiteri's model, Xiao et al.<sup>21</sup> use Lennard-Jones (LJ) potentials rather than Buckingham potentials to describe all van der Waals interactions. They included a LJ potential for the C\_C interaction, which was neglected in Pavese's work<sup>18</sup>.

The simulation model of surface follows the approach of Tasker<sup>42</sup> in which the crystal consists of two blocks, each comprising two regions that are periodic in two dimensions. Region I contains the surface layers and few layers immediately below; these atoms are allowed to relax to their mechanical equilibrium. Region II contains those atoms further away, which are kept fixed at their bulk equilibrium position and represent the rest of the crystal. The size of region 2 is chosen to be thick enough so that the atoms represent the effect of the bulk upon the atoms in region 1. The size of region 1 is generally chosen depending upon the convergence of surface energy. This quantity is computed using the following expression:

$$\gamma = \frac{U_s - U_b}{A}$$

where  $U_s$  is the energy of the surface block of the crystal,  $U_b$  is the energy of an equal number of atoms of the bulk crystal, and A is the surface area. The energies of the blocks are essentially the sum of the energies of interaction between all atoms. The surface energy is a measure of the thermodynamic stability of a cleavage plane. On the





(b)

(10-1)

(0-10) (1-1-1)

Figure 8 | Equilibrium morphology of hydrous CaCO<sub>3</sub> polymorphs: (a) Ikaite compared to SEM image<sup>39</sup>; (b) monohydrocalcite.

other hand there is a widely used alternative criterion named the attachment energy: it is the energy associated with the addition of a stoichiometric layer of material onto the surface cut. This implies that for any stable material, the attachment energy must be exothermic. Therefore, those faces will tend to grow most rapidly.

- 1. Morse, J. W. & Mackenzie, F. T. Geochemistry of Sedimentary Carbonates. Develop. Sedimentol. 48, Elsevier: Amsterdam, 724 pages (1990).
- Usher, C. R. et al. Reactions on mineral dust. Chem. Rev. 103, 4883-4939 (2003).
- Morse, J. W. et al. Calcium carbonate formation and dissolution. Chem. Rev. 107, 342-381 (2007).
- Van Cappellen, P. et al. A surface complexation model of the carbonate mineralaqueous solution interface. Geochim. Cosmochim. Acta. 57, 3505-3518 (1993).
- Swainson, I. P. The structure of monohydrocalcite and the phase composition of the beachrock deposits of Lake Butler and Lake Fellmongery. South Australia 93,
- Hull, H. & Turnbull, A. G. Thermochemical Study of Monohydrocalcite. Geochim. Cosmochim. Acta 37, 685-694 (1973).
- Dahl, K. & Buchardt, B. Monohydrocalcite in the arctic Ikka fjord, SW Greenland: First reported marine occurrence. Jour. of Sedim. Res. 76, 460-471 (2006).
- Lennie, A. R. et al. The structure and thermal expansion behaviour of ikaite,  $CaCO_36H_2O$ , from T = 114 to T = 293 K. Miner. Mag. 68, 135-146 (2004).
- Bischoff, J. L. et al. The Solubility and Stabilization of Ikaite (CaCO<sub>3</sub>.6H<sub>2</sub>O) from 0° to 25°C: Environmental and paleoclimatic implications for thinolite tufa. The Jour. of Geo. 101, 21-33 (1993).
- 10. Marland, G. The stability of CaCO3.6H2O (ikaite). Geoch. et Cosmoch. Acta 39, 83-91 (1975).
- 11. Shahar, A. et al. The stability and Raman spectra of ikaite, CaCO<sub>2</sub>.6H<sub>2</sub>O, at high pressure and temperature. Amer. Miner. 90, 1835-1839 (2005).
- 12. Jansen, J. H. F. et al. Ikaite pseudomorphs in the Zaire deep-sea fan: An intermediate between calcite and porous calcite. Geol. 15, 245-248 (1987).
- 13. Frye, R. editor. Aragonite Group in Encyclopedia of Mineralogy. Hutchinson Ross Publishing Company, Stroudsbourg, Pennsylvania, USA, 1 edition (1981).
- 14. De Leeuw, N. H. & Parker, S. C. Surface structure and morphology of calcium carbonate polymorphs calcite, aragonite, and vaterite: An atomistic approach. Jour. of Phys. & Chem. B 102, 2914-2922 (1998).
- 15. De Leeuw, N. H. & Parker, S. C. Atomistic simulation of the effect of molecular adsorption of water on the surface structure and energies of calcite surfaces. J. Chem. Soc. Faraday Trans. 93(3), 467-475 (1997).
- 16. Pavese, A. et al. Interatomic potentials for CaCO3 polymorphs (calcite and aragonite), fitted to elastic and vibrational data. Phys. & Chem. of Miner. 19, 80-87

- 17. Archer, T. D. et al. An interactomic potential model for carbonates allowing for polarisation effects. Phys. & Chem. of Miner. 30, 416-424 (2003).
- 18. Pavese, A. et al. Modelling of the thermal dependence of structural and elastic properties of calcite, CaCO<sub>3</sub>. Phys. & Chem. of Miner. 23, 89-93 (1996).
- 19. Raiteri, P. & Gale, J. D. Water Is the Key to Nonclassical Nucleation of Amorphous Calcium Carbonate. Amer. Chem. Soc. 132, 17623-17634 (2010).
- 20. Raiteri, P. et al. Derivation of an Accurate Force-Field for Simulating the Growth of Calcium Carbonate from Aqueous Solution: A New model for the Calcite-Water Interface. Jour. of Phys. & Chem. C. 114, 5997-6010 (2010).
- 21. Xiao, S. et al. A New Transferable Forcefield for Simulating the Mechanics of CaCO<sub>3</sub> Crystals. Jour. of Phys. Chem. C. 115, 20067-20075 (2011).
- 22. Fisler, D. K. et al. A shell model for the simulation of rhombohedral carbonate minerals and their point defects. Amer. Miner. 85, 217-224 (2000).
- 23. Sekkal, W. et al. A lattice dynamical study of the aragonite and post-aragonite phases of calcium carbonate rock. Amer. Miner. 93, 1608-1612 (2008).
- 24. Dove, M. T. et al. A new interatomic potential mlodel for calcite: Applications to lattice dynamics studies, phase transition, and isotope fractionation. Amer. Miner. 77, 244-250 (1992).
- 25. Catti, M. et al. Thermodynamic properties of CaCO<sub>3</sub> calcite and aragonite: A quasi-harmonic calculation. Phys. & Chem. of Miner. 19, 472–479 (1993).
- 26. Parker, S. C. et al. Phil. Trans. of the Roy. Soc. of London Series A Phys. Sci. and Eng. 344, 37 (1993).
- 27. Gale, D. J. GULP: A computer program for the symmetry-adapted simulation of solids. Chem. Soc. Faraday Transactions. 93, 629-637 (1997).
- 28. Glasser, L. & Jenkins, H. D. B. Lattice energies and unit cell volumes of complex ionic solids. J. Amer. Chem. Soc. 122, 632-638 (2000).
- 29. Chen, C. C. et al. Elasticity of single-crystal calcite and rhodochrosite by Brillouin spectroscopy. Amer. Miner. 86, 1525-1529 (2001).
- 30. Liu, L. et al. Elasticity of single-crystal aragonite by Brillouin spectroscopy. Phys. and Chem. of Miner. 32, 97-102 (2005).
- 31. Hill, R. Elastic properties of reinforced solids: some theoretical principles. Jour. of the Mech. and Phys. of Solids. 11, 357-372 (1963).
- 32. Pugh, S. F. Relations between the elastic moduli and the plastic properties of polycrystalline pure metals. Philos. Mag. 45, 823-843 (1954).
- 33. Ayoub, A. et al. High pressure structural phase transitions and mechanical properties of calcite rock. Comput. Mater. Sci. 50, 852-857 (2011).
- 34. Dove, P. M. & Hochella, M. F. Calcite precipitation mechanisms and inhibition by orthophosphate: In situ observations by scanning force microscopy. Geoch. Cosmoch. Acta 57, 705-714 (1993).
- 35. Heywood, B. R. & Mann, S. Molecular Construction of Oriented Inorganic Materials: Controlled Nucleation of Calcite and Aragonite under Compressed Langmuir Monolayers. Chem. of Mater. 6, 311 (1994).



- 36. Wulff, G. Zur Frage der Geschwindigkeit des Wachstums und der Auflösung der Krystallflagen. Zeitsch. für Kristall. 34, 449–480 (1901).
- 37. Gibbs, J. W. Collected Works. Longma. New York (1928).
- 38. Mukkamala, S. B. *et al.* Modelling calcium carbonate biomineralisation processes. *Jour. of Inorg. Biochem.* **100**, 1128–1138 (2006).
- Dieckmann, G. S. et al. Ikaite (CaCO<sub>3</sub>\*6H<sub>2</sub>O) discovered in Arctic sea ice. The Cryosphere, 4, 227–230 (2010).
- Wu, Y. J. et al. Flexible simple point-charge water model with improved liquidstate properties. J. of Chem. Phys. 124, 024503 (2006).
- 41. Jorgensen, W. L. *et al.* Comparison of simple potential functions for simulating liquid water. *The J. of Chem. Phys.* **79**, 926–935 (1983).
- 42. Tasker, P. W. The stability of ionic crystal surfaces. *J. of Phys. C. Solid State Phys.* 12, 4977–4983 (1979).

# **Acknowledgments**

The authors would like to thank the National Radioactive Waste Management Agency (ANDRA) for financial support. The authors are grateful to Dr F. Plas and A. Pasteau for valuable discussions.

## **Author contributions**

A.Z. conceived and designed the study. W.S. performed the theoretical studies and simulations. W.S. and A.Z. analysed and discussed the results. W.S. and A.Z. wrote the manuscript. All authors discussed the conceptual and practical implications of the method at all stages.

### Additional information

Competing financial interests: The authors declare no competing financial interests.

**License:** This work is licensed under a Creative Commons Attribution-NonCommercial-NoDerivs 3.0 Unported License. To view a copy of this

license, visit http://creativecommons.org/licenses/by-nc-nd/3.0/ **How to cite this article:** Sekkal, W. & Zaoui, A. Nanoscale analysis of the morphology and surface stability of calcium carbonate polymorphs. *Sci. Rep.* 3, 1587; DOI:10.1038/ srep01587 (2013).