
Updated seismotectonic framework of Abu Dabbab Egypt based on focal mechanisms and stress inversion

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1 **Updated Seismotectonic Framework of Abu Dabbab Egypt based on Focal
2 Mechanisms and Stress Inversion**

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16 **Abstract**

17 The Abu Dabbab seismic zone is located along Egypt's Red Sea margin,
18 stands out as one of the most active seismic regions in the Eastern Desert.
19 Characterized by frequent micro earthquakes, swarm like activity, and notable
20 historical events. To enhance understanding of its tectonic framework, 408
21 earthquakes (M1 0.7–3.0) recorded in 2004 were analyzed using digital waveform
22 data from a temporary local seismic network consisting of ten vertical short
23 period seismometers. Focal mechanisms were determined from P-wave first-
24 motion polarities and classified using ternary plots. The analysis revealed a
25 diverse range of faulting styles normal, strike slip, reverse, and oblique with
26 clear depth dependent patterns. Shallow events (0–5 km) were dominated by
27 normal and strike slip faulting, intermediate depths (5–10 km) showed increased
28 reverse and oblique components, while deeper events (>10 km) were primarily
29 normal faulting. Stress tensor inversion across three depth intervals indicated a
30 multiphase stress regime: shallow depths exhibited alternating faulting styles
31 due to localized stress variations; intermediate depths revealed a heterogeneous
32 stress field with mixed faulting regimes; and deeper levels showed a dominant
33 normal faulting regime, consistent with the extensional tectonics of the Red Sea
34 Rift. Overall, the stress field is shaped by NE-SW compression and SE-NW
35 extension, with deformation concentrated along NW-SE and NE-SW trending
36 faults. These findings underscore the combined influence of regional rift-related
37 extension and local factors such as magmatic intrusions and crustal
38 heterogeneity in driving seismicity at Abu Dabbab. This study yields important

39 insights into depth dependent stress patterns and active faulting, enhancing
 40 seismic hazard assessments and highlighting the region's potential as a
 41 sustainable geothermal energy source within a tectonically dynamic
 42 environment.

43

44 **Keywords:** Seismicity, Focal mechanism, stress tensor, Abu Dabab, Egypt

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49 1. Introduction

50 Abu Dabbab is situated along Egypt's Red Sea coast, approximately 24
 51 kilometers inland and 30 kilometers north of Marsa Alam (25.15°–25.35°N,
 52 34.35°–34.65°E) it is located within the tectonically active Eastern Desert, it
 53 forms part of the Red Sea Rift System a segment of the Afro Arabian Rift shaped
 54 by the divergent motion of the African and Arabian plates. This region is notable
 55 for recurrent micro earthquake swarms, such as those recorded in 1984 and
 56 between 2004 and 2005, which occurred without a distinct mainshock. The name
 57 "Abu Dabbab" itself carries seismic significance, translating to "earthquake
 58 sounds" (from Dabbab meaning knocking, and Abu meaning father) a
 59 phenomenon long described by local Bedouin communities who reported hearing
 60 these mysterious subterranean sounds. Strategically positioned at the
 61 intersection of major tectonic and structural trends, Abu Dabbab is recognized as
 62 one of Egypt's most seismically active zones.

63

64 The region has long intrigued researchers due to its persistent and
 65 anomalous seismic activity. Early studies by [1-3] documented major events such
 66 as the 1955 (mb=6) and 1984 (M=5.1) earthquakes and proposed that the
 67 seismicity might be related to a plutonic intrusion within the Precambrian crust,
 68 rather than solely regional tectonic forces. This view was supported by the tight
 69 clustering of hypocenters and sustained swarm-like activity. Additional studies
 70 have linked the region's seismic behavior to igneous intrusions and potential
 71 magma sources, citing its unusually high heat flow (~92 mW/m²), nearly double
 72 the national average [4,5] and seismic tomography revealing P and S wave
 73 anomalies indicative of magma intrusion beneath the area [6] while observations
 74 of co seismic surface deformation near earthquake source zones point to a
 75 heterogeneous crustal structure [7,8]. These findings suggest a complex interplay
 76 between tectonic stress and magmatic processes in driving Abu Dabbab's seismic
 77 dynamics. Despite these findings, debate continues, with some researchers

78 suggesting that both tectonic stress and magmatic processes contribute to the
 79 area's seismic behavior [9,10]. Abu Dabbab's tectonic setting is intimately
 80 connected to its persistent seismic activity, as documented by numerous
 81 researchers [11,12]. In addition to the major earthquakes of 1955 and 1984, the
 82 region has experienced recurring seismic swarms, frequently accompanied by
 83 audible rumbling sounds reported by local communities [5]. Notable swarms were
 84 recorded in 1976, 1984, and 1993 [6, 9,11].

85

86 The faulting mechanisms in Abu Dabbab have evolved over time, reflecting
 87 the region's complex tectonic dynamics. The 1955 earthquake analysis by [13]
 88 revealed a strike slip faulting mechanism with a normal dip slip component, with
 89 nodal planes trending NNW-NW and ENE-ESE. In contrast, the 1984 event
 90 exhibited strike slip faulting with a normal component [10]. During the August
 91 2004 seismic swarm, [9] documented a shift in faulting style: while the
 92 background seismicity was dominated by normal faulting with strike slip
 93 elements, the swarm itself was characterized by reverse faulting with strike slip
 94 components. Consistent fault orientations NW, NE, ENE-WNW, and NNE-NNW
 95 were interpreted as indicative of left lateral strike slip faulting. [10] further
 96 identified thrust faulting mechanisms during the 2004 swarm, contrasting with
 97 earlier events such as the 1984 earthquake, which showed strike slip faulting
 98 with a normal component. These findings suggest right lateral slip along NE-SW
 99 faults and left lateral slip along NW-SE faults, with T-axis trends ranging from
 100 NNE-SSW to NNW-SSE. This pattern reflects the reactivation of structures
 101 associated with the Najd Fault System. Subsequent studies confirmed the
 102 coexistence of normal, reverse, and strike-slip faulting styles in the region [14-16].
 103 More recently, [17] reported NE-SW compression and SE-NW extension,
 104 consistent with broader regional strain patterns.

105 The Abu Dabbab region hosts extensive geological formations with ultra
 106 basic to basic intrusions, including flows, dikes, and massive bodies. It is defined
 107 by three major wadis Abu Dabbab, Mubarak, and Dabr. Wadi Mubarak, in the
 108 north, contains thick sequences of low grade volcanic sedimentary rocks of back
 109 arc and arc affinities, intruded by both older and younger granitic bodies
 110 [18,19]. The broader Eastern Desert of Egypt, particularly along the Red Sea
 111 margin, holds significant geothermal energy potential, supporting Egypt's
 112 strategy to expand sustainable energy resources in tectonically active areas.

113 In regions with geological and tectonic settings comparable to Abu
 114 Dabbab such as the Gediz Graben in western Anatolia researchers have employed
 115 diverse geophysical techniques to investigate crustal processes. These areas are
 116 characterized by active extensional tectonics, pronounced lateral variations in
 117 sedimentary cover thickness, and significant magmatic activity [20, [21] advanced

118 this work by integrating continuous, high resolution seismic monitoring with
 119 complementary geophysical surveys, including magneto telluric and state of the
 120 art tomography. This combined methodology enables precise imaging of
 121 intrusion geometries, monitoring of stress evolution, and direct evaluation of
 122 geothermal systems. Such integration is essential for unraveling the dynamics of
 123 active rift margins and plays a critical role in reducing exploration risks for
 124 renewable energy projects in tectonically complex environments.

125 Although seismicity in Abu Dabbab has been widely studied, detailed stress
 126 tensor inversion analyses are limited. Earlier work revealed a clockwise rotation
 127 of the T-axis and a complex strain regime, highlighting the interplay of
 128 compressional and extensional forces.

129 This study advances the field by analyzing focal mechanism solutions for
 130 408 micro earthquakes (M_l 0.7–3.0) recorded in 2004, enabling a comprehensive,
 131 depth dependent stress tensor inversion. Results show deformation patterns
 132 consistent with regional tectonic stress orientations, with seismicity driven by
 133 stress concentrations near the tip of a propagating intrusion. The findings
 134 improve understanding of depth dependent stress regimes and active faulting,
 135 strengthen seismic hazard assessments, and emphasize Abu Dabbab's potential
 136 as a sustainable geothermal energy resource in a tectonically dynamic setting.

137 2. Data Source and Methodology

138 The present study utilizes raw waveform data recorded in 2004 by ten
 139 temporary seismic stations deployed in the Abu Dabbab region by the Egyptian
 140 National Seismological Network (ENSN). Earthquakes were extracted from the
 141 raw data and located using the HYPOINVERSE software, based on the crustal
 142 velocity model of [22]. To better understand the tectonic framework of the study
 143 area, the analysis was conducted separately for each month. In 2004, an intense
 144 earthquake swarm occurred, producing more than 4,000 recorded events within
 145 just a few months. From this sequence, 408 earthquakes were selected for
 146 detailed examination in order to identify any apparent changes in focal
 147 mechanisms and stress tensor during the swarm period.

148 Fault plane solutions were determined for 408 earthquakes with local
 149 magnitudes ranging from 0.7 to 3.0 using the classical first motion polarity
 150 method of P waves by PMAN software [23] to obtain solutions depending on the
 151 azimuth, incidence angle, and polarities of P-phase only. Furthermore, [22]
 152 demonstrates the importance of high precision earthquake relocation and cluster
 153 analysis in deciphering seismogenic processes.

154 QP was calculated for all focal mechanism solutions where QP is the quality
 155 code that describes the range of uncertainties in the strike, dip, and rake of a

156 focal mechanism solution. To approach this purpose we should obtain acceptable
 157 solutions for each event thus we used the FOCMEC software [24], which performs
 158 a grid search over all possible solutions based on user selected parameters,
 159 including polarity errors and the allowable disagreement between observed and
 160 calculated solutions. In this study, solutions were estimated using a 5° grid
 161 search. QP and azimuth gap for each event are illustrated in supplementary
 162 material.

163 In this study, we applied the stress tensor inversion technique of [25] to
 164 characterize the stress field in the Abu Dabbab region. Stress tensor inversion
 165 methods, originally developed by [26-29] and later refined by [29, 30], assume that:
 166 The regional tectonic stress is homogeneous, Earthquakes occur on pre-existing
 167 faults with varying orientations, and Slip on a fault plane occurs in the direction
 168 of maximum shear stress (τ), according to the Wallacem Bott hypothesis [31,32].
 169 [32] defined four independent parameters describing the reduced stress tensor:
 170 the principal stresses σ_1 , σ_2 , σ_3 , and the stress ratio R , which governs the
 171 orientation of shear stress on fault planes. We used TENSOR software [25] to
 172 estimate these parameters. Focal mechanisms incompatible with the
 173 predominant stress field are filtered out. Horizontal stress orientations (SHmax
 174 and SH_{min}) are computed following [33]. Recent advances, such as the framework
 175 proposed by [34,35], emphasize the integration of multi depth and multi-fault
 176 kinematic data for stress tensor inversion, particularly for hidden seismogenic
 177 faults. Their methodology highlights the necessity of detailed fault plane
 178 solutions and stress ratio assessments in complex tectonic settings, which aligns
 179 closely with our depth-dependent stress regime analysis in Abu Dabbab.

180 The quality of the stress inversion results is evaluated using the World
 181 Stress Map (WSM) ranking system [25,36]:

- 182 □ A-quality: $N \geq 15$ events, $\alpha \leq 12^\circ$, SHmax/S_{Hmin} $\pm 15^\circ$
- 183 □ B-quality: $8 < N < 15$, $12^\circ < \alpha \leq 20^\circ$, SHmax/S_{Hmin} $\pm 15-20^\circ$
- 184 □ C-quality: $6 \leq N < 8$ or $\alpha > 20^\circ$, SHmax/S_{Hmin} $\pm 20-25^\circ$
- 185 □ D-quality: Only 4-5 events per box, SHmax/S_{Hmin} $\pm 25-40^\circ$

186 This approach allows robust, depth dependent characterization of the stress field
 187 based on the compiled focal mechanism solutions from the Abu Dabbab region.

188

189 **3. Seismicity in Abu Dabbab Seismic Zone**

190 The Abu Dabbab region in Egypt's Eastern Desert has long intrigued
 191 researchers due to its persistent and anomalous seismic activity. [1], using
 192 WWSSN data to analyze Red Sea seismicity since 1953, reported a major
 193 earthquake on 12 November 1955 (mb = 6), followed by a moderate event on 2

194 July 1984 ($M = 5.1$). ^[2,3] attributed this seismicity to a plutonic intrusion within
 195 the Precambrian crust, rather than to regional tectonic forces a view supported
 196 by the tight clustering of hypocenters and the sustained activity over time.
 197 Several studies have linked seismic activity in Abu Dabbab to igneous intrusions
 198 and underlying magma sources, citing the region's exceptionally high heat flow
 199 (approximately 92 mW/m^2 , nearly twice the national average) ^[4,5]. Despite this
 200 thermal evidence, the origin of the seismicity remains a topic of ongoing debate.
 201 Some researchers argue that both tectonic stress and magmatic processes jointly
 202 contribute to the area's complex seismic behavior ^[9,10].

203 Seismicity maps produced in this study (Fig. 1A & 1B), based on the
 204 relocation of 2,333 earthquakes recorded in 2004 (histogram in Fig. 2), reveal a
 205 clear NE-SW alignment of seismic events. This trend is nearly perpendicular to
 206 the Red Sea axis and coincides with a major fault system (Najd Fault System).
 207 The majority of relocated earthquakes are concentrated in the southern part of
 208 the region, occurring at shallow depths between 0 and 20 km, with magnitudes
 209 ranging from 0.7 to 3.8. These events were detected by ten temporary seismic
 210 stations deployed during the study. The relocation results demonstrate high
 211 accuracy, with RMS residuals ranging from 0.01 to 0.4, horizontal errors
 212 between 0.02 and 2 km, and vertical depth errors between 0.05 and 1.0 km as
 213 shown in Fig. (3). The predominance of swarm like seismicity highlights the
 214 unique geological characteristics of Abu Dabbab, setting it apart from other
 215 seismically active regions in Egypt.

216 4. Focal Mechanism Solution in Abu Dabbab Seismic Zone

217 In this study, fault plane solutions were determined for 408 earthquakes in
 218 the Abu Dabbab region, with local magnitudes (M_l) ranging from 0.7 to 3.0.
 219 Digital waveform data were acquired from a temporary local seismic network
 220 installed in the area. Focal mechanisms were constructed using P-wave first
 221 motion polarities, with events carefully relocated to enhance the accuracy of
 222 azimuth, incidence angle, and polarity readings. This refinement enabled more
 223 reliable use of key waveform parameters essential for fault-plane analysis. The
 224 classical polarity-based method was applied using PMAN software ^[23] to derive
 225 the fault plane solutions.

226 To visualize faulting styles, we applied ^[37], which plot focal mechanisms
 227 within a triangular framework defined by three end member fault types: pure
 228 strike-slip, normal, and thrust faulting ^[37-39]. This approach provides a clear and
 229 intuitive framework for comparing the relative distribution of faulting styles
 230 within the dataset. Since earthquake activity was concentrated between May and
 231 July, focal mechanism solutions were developed independently for each month to

232 capture temporal variations. Cross sectional analysis revealed that seismic
 233 events were distributed across three main depth intervals: 0–5 km, 5–10 km, and
 234 10–20 km, offering insights into the vertical segmentation of faulting activity
 235 (Fig. 4A, 4B &4C).

236

237 The analysis of 408 focal mechanism solutions reveals a wide range of
 238 faulting styles in the Abu Dabbab region, including normal, strike slip, reverse,
 239 and oblique faulting. These variations are effectively illustrated using ternary
 240 plots (Fig. 5A, 5B &5C). Example of hypocentral parameters for selected events
 241 are presented in Table 1, with corresponding focal mechanism solutions in Table
 242 2. The complete hypocentral dataset is provided in the supplementary material.
 243 Figs. 6, 7, and 8 illustrate the spatial distribution of seismic events across the
 244 Abu Dabbab region, accompanied by focal mechanism "beach ball" diagrams and
 245 detailed beach balls with stations distribution are provided in the supplementary
 246 material.

247 **4.1 Results of focal mechanism solutions of earthquakes in Abu**
 248 **Dabbab:**

249 Comparative analysis across depth intervals reveals distinct variations in
 250 faulting styles, reflecting the region's complex and layered tectonic structure. In
 251 the shallow depth range (0–5 km), most earthquakes exhibit normal, normal
 252 oblique, and strike slip mechanisms, with reverse faulting playing only a minor
 253 role. At intermediate depths (5–10 km), faulting styles shift noticeably, with
 254 reverse and oblique mechanisms becoming more dominant and a decline in
 255 strike slip activity. Beyond 10 km, faulting styles are more evenly distributed
 256 among all categories, indicating a broader range of tectonic processes at greater
 257 depths. Overall, the Abu Dabbab region is shaped by NW-SE and NE-SW
 258 trending fault systems, with strike slip components exhibiting depth dependent
 259 variability. The constructed focal mechanism solutions serve as the foundation
 260 for a comprehensive stress tensor inversion, offering valuable insights into the
 261 region's active deformation and prevailing stress regimes. These results are
 262 consistent with the broader tectonic stress orientation and suggest that the
 263 observed seismicity is driven by stress concentrations near the tip of a
 264 propagating magmatic intrusion [15,40,41].

265 **5. Stress Tensor Inversion in Abu Dabbab Region**

266 To assess the stress field in the Abu Dabbab region, we applied stress
 267 tensor inversion to a dataset of 408 focal mechanism solutions. The inversion

268 was performed using the method developed by [24] implemented through the
 269 TENSOR software. Stress inversion techniques are widely recognized for their
 270 effectiveness in estimating regional stress fields from earthquake focal
 271 mechanisms [26-29]. As mentioned before, the reduced stress tensor is defined by
 272 four key parameters: the orientations of the principal stress axes (σ_1 , σ_2 , σ_3) and
 273 the stress ratio (R), which determines the geometry of fault slip [32]. In this study,
 274 the dataset was divided into three depth intervals, consistent with the
 275 classification used in the focal mechanism analysis, to explore depth dependent
 276 variations in the regional stress field.

277

278 **5.1 Results of Stress Tensor Inversion in Abu Dabbab**

279 **5.1.1 At Shallow Depth (0-5 Km)**

280 □ Based on the stress tensor inversion model developed for
 281 the Abu Dabbab region, three distinct stress regimes were identified
 282 (Fig. 9):

283 □ Normal Faulting Regime (Fig. 9A): Characterized by a nearly
 284 vertical σ_1 (plunge 88°) and a horizontal σ_3 (plunge 0°), with a stress
 285 ratio R' equal to R. This configuration indicates an extensional stress
 286 field, with the minimum horizontal stress (Sh_{min}) oriented N75°E. The
 287 solution quality is high (Grade A), supported by a low misfit angle (α =
 288 10.3) and an average misfit function value of $F5 = 5.1$ (Table 3).

289 □ Oblique (Normal Strike-Slip) Regime (Fig. 9B): Features moderately
 290 dipping σ_1 and σ_2 (plunges of 49° and 40°, respectively), with a
 291 horizontal σ_3 (plunge 6°), and a stress ratio of $R' = 1.0$. This
 292 configuration indicates an extensional stress field with a strike slip
 293 component, where the minimum horizontal stress (Sh_{min}) trends N78°E.
 294 The solution is of high quality (Grade A), supported by a low misfit angle
 295 ($\alpha = 8.4$) and an average misfit function value of $F5 = 3.5$ (Table 3).

296

297 □ Thrust Faulting Regime (Fig. 9C): Characterized by a horizontal
 298 maximum stress axis (σ_1) and a nearly vertical minimum stress axis (σ_3),
 299 indicating a compressional stress field. The stress ratio is $R' = 2.5$, and
 300 the solution is of high quality (Grade A), with a misfit angle of 10.1 and
 301 an average misfit function of 5.6 (Table 3).

302 **5.1.2 At Intermediate Depth (5-10 Km)**

303 Stress tensor inversion for the region revealed four distinct stress types, as
 304 shown in Figure 10:

- 305 □ Normal Faulting (Fig. 10A): Characterized by a sub-vertical
 306 maximum principal stress axis (σ_1 , plunge 78°) and a horizontal
 307 minimum principal stress axis (σ_3 , plunge 1°), with a stress ratio of $R' =$
 308 R . This configuration indicates an extensional stress field, with the
 309 minimum horizontal stress (Sh_{min}) trending N86°E. The solution is of
 310 high quality (Grade A), supported by a misfit angle ($\alpha = 11.4$) and an
 311 average misfit function value of $F5 = 5.9$ (Table 3).
- 312 □ Oblique (Normal Strike-Slip) Regime (Fig. 10B): Features
 313 moderately dipping σ_1 and σ_2 (plunges of 53° and 37°, respectively), with
 314 a horizontal σ_3 (plunge 3°), and a stress ratio of $R' = 1$. This
 315 configuration indicates an extensional stress field with a strike slip
 316 component, where the minimum horizontal stress (Sh_{min}) trends N75°E.
 317 The solution is of high quality (Grade A), supported by a low misfit angle
 318 ($\alpha = 9.7$) and an average misfit function value of $F5 = 4.5$ (Table 3).
- 319 □ Strike-Slip Regime (Fig. 10C): In this regime, the maximum (σ_1) and
 320 minimum (σ_3) principal stresses are both sub-horizontal (with plunges of
 321 06° and 01°, respectively), while the intermediate principal stress (σ_2) is
 322 nearly vertical (Table 3, Fig. 9C). The related stress tensor is of B
 323 quality, featuring a low misfit angle (α) of 7.9° and an average misfit
 324 function ($F5$) of 3.2. This regime is characterized by an extensional
 325 direction of N60°E and a stress ratio (R') of 1.5.
- 326 □ Thrust Faulting Regime (Fig. 10D): Characterized by a horizontal
 327 maximum principal stress axis (σ_1 , plunge 0°), a nearly vertical minimum
 328 stress axis (σ_3 , plunge 73°), and a sub-horizontal intermediate axis (σ_2).
 329 This configuration indicates a compressional stress field, with a stress
 330 ratio of $R' = 2.50$. The solution is rated high quality (Grade A), supported
 331 by a misfit angle of 10.6 and an average misfit function value of $F5 = 4.8$
 332 (Table 3).

333 **5.1.3 At Depth Greater than 10 Km**

334 For seismic events deeper than 10 km, the optimal stress tensor inversion
 335 indicates an oblique tectonic regime characterized by normal and strike slip
 336 faulting components (Fig. 11). This regime features a sub vertical σ_1 (plunge 55°)
 337 and a horizontal σ_3 (plunge 1°), with a stress ratio $R' = 1$. The resulting oblique
 338 normal faulting is oriented with a minimum horizontal stress (Sh_{min}) direction of

339 N77°E. This high quality, Grade A solution is constrained by a low misfit angle (α
 340 = 9°) and an average misfit function value (F5) of 3.8 (Table 3).

341 The results demonstrate a complex stress history in the Abu Dabbab region,
 342 with evidence for normal, strike slip, thrust, and oblique faulting regimes across
 343 different depths. Normal faulting dominates at shallow and deep levels, while
 344 intermediate depths reveal a more varied stress field, reflecting transitional
 345 tectonic processes. This pattern can be interpreted in the context of dike
 346 intrusion induced stresses: compression in the walls of a dike and tension above
 347 a propagating dike tip, with hypocenters potentially migrating ahead of the
 348 advancing dike front.

349 **6. Discussion**

350 This research delivers a depth resolved analysis of stress and deformation
 351 within the Abu Dabbab seismic zone. By integrating focal mechanism solutions
 352 with stress tensor inversion across multiple crustal levels, the study advances
 353 beyond simple tectonic classification. The findings reveal a vertically layered
 354 seismogenic system governed by the combined influence of regional tectonic
 355 forces and localized magmatic processes.

356 The key finding of this study is the distinct vertical partitioning of faulting
 357 styles and stress regimes, which together define a coherent mechanical model of
 358 crustal deformation beneath Abu Dabbab. In the shallow crust (0–5 km),
 359 deformation is dominated by normal and strike slip faulting under normal and
 360 oblique normal stress regimes. This pattern reflects the regional Red Sea
 361 extensional tectonics, driven by SE–NW extension. However, the recurring
 362 presence of a localized thrust stress regime at these depths indicates a notable
 363 disturbance within the regional stress field.

364 The shallow compressional pocket is best explained by stress rotation or
 365 clamping, resulting from the mechanical interaction between the extending crust
 366 and a shallow, rigid intrusive body, or alternatively from the complex geometry
 367 of intersecting and reactivated Najd fault systems. At depths of 5–10 km, the
 368 crust becomes a pronounced stress transition zone marked by strong
 369 heterogeneity. Within this interval, all four modeled stress regimes normal,
 370 oblique, strike slip, and thrust are present, indicating that the crust is subject to
 371 competing and spatially variable forces rather than a uniform stress field. This
 372 depth range is interpreted as the key interaction zone where regional
 373 extensional stresses from the Red Sea rift intersect with localized perturbations
 374 driven by an actively pressurizing magmatic intrusion. The lateral and vertical
 375 migration of magma can impose compressional stresses along intrusion flanks
 376 and generate transient compression ahead of its advancing tip (Fig.12), while

377 regional extension continues to dominate the surrounding crust. The coexistence
 378 of reverse and normal faulting within this interval a feature widely reported in
 379 magmatically active rift settings serves as a diagnostic marker of dynamic
 380 tectonic magmatic interplay. At depths greater than 10 km, the stress field
 381 becomes more homogeneous, with deformation primarily governed by oblique
 382 normal faulting.

383 This pattern indicates that the influence of local magmatic perturbations
 384 and near surface crustal heterogeneity diminishes with depth, leaving
 385 deformation at greater levels primarily governed by a stable regional extensional
 386 stress regime. Such deep seated extension is consistent with large scale crustal
 387 necking beneath the Red Sea margin and likely represents the fundamental
 388 tectonic driver into which the magmatic system is emplaced.

389 Our depth dependent analysis strongly supports a hybrid model that
 390 resolves the longstanding debate over the source of Abu Dabbab's seismicity.
 391 Regional NE-SW compression and SE-NW extension establish the tectonic
 392 framework, orienting the principal stress axes and favoring slip along two
 393 dominant fault sets: NW-SE trending faults (normal to oblique slip) and NE-SW
 394 trending faults (strike slip).

395 This explains the consistent fault orientations and the broader regional
 396 kinematic framework. However, the anomalous features such as intense swarm
 397 like clustering, elevated heat flow ($\sim 92 \text{ mW/m}^2$), the coexistence of contradictory
 398 stress regimes in close proximity, and reports of seismic sounds cannot be
 399 attributed to regional tectonics alone. Our model proposes that an actively
 400 pressurizing magmatic intrusion is the key local driver of Abu Dabbab's
 401 seismicity. Acting as a stress concentrator, it amplifies and rotates stresses,
 402 producing the heterogeneous regimes observed. The seismicity pattern
 403 particularly the migration and swarm behavior noted in earlier studies is
 404 interpreted as the response to dike propagation: hypocenters advance with the
 405 dike tip under tensile stresses, while compressive events occur along its flanks.
 406 In this framework, regional extension establishes the fractured crust and
 407 fundamental stress orientation, while the magmatic intrusion supplies the
 408 dynamic energy and localized stress complexity that sustain swarm activity. This
 409 interpretation is supported by seismic tomography studies showing velocity
 410 anomalies ^[6] and is consistent with global analogues of seismicity in
 411 magmatically active rift systems. Recognizing a multiphase, depth dependent
 412 stress regime has significant implications for seismic hazard assessment in the
 413 Abu Dabbab region. Faults are not limited to a single kinematic style; instead,
 414 the same NE-SW or NW-SE trending structures may shift between normal,
 415 strike-slip, or reverse motion depending on transient local stress perturbations

416 and depth. This variability increases uncertainty in defining seismic source
 417 mechanisms and complicates the development of fault based hazard models.

418 The presence of localized compressional stress regimes indicates that the
 419 crust can accumulate and release strain through reverse faulting, meaning
 420 hazard assessments must account for the possibility of moderate magnitude
 421 earthquakes ($M \sim 5-6$) driven by non extensional mechanisms. Such events may
 422 generate ground motion characteristics distinct from those expected in typical
 423 normal faulting earthquakes within extensional settings. Furthermore, the
 424 spatial clustering of heterogeneous seismicity and swarm activity in the southern
 425 sector particularly within the 5–10 km depth range marks this zone as an area of
 426 elevated seismic potential that should be prioritized for targeted monitoring and
 427 hazard mitigation efforts. The tectonic magmatic framework developed in this
 428 study significantly strengthens and de-risks the geothermal potential of the Abu
 429 Dabbab region. Seismicity is interpreted as direct evidence of an active mid
 430 crustal magmatic heat source, providing a robust explanation for the
 431 exceptionally high surface heat flow reported in the area [4,5]. The predominance
 432 of normal and oblique faulting within both shallow and deep extensional regimes
 433 further indicates that the crust is extensively fractured by ongoing extension.
 434 This interconnected fault and fracture network enhances permeability, enabling
 435 efficient fluid circulation and heat extraction. Moreover, the persistence of
 436 seismic activity over several decades points to a long lived system of heat and
 437 fluid transport, a defining characteristic of sustainable geothermal resources.

438 Collectively, these findings establish Abu Dabbab as a scientifically
 439 validated, high priority target for geothermal exploration and development. The
 440 study reframes the area from being simply seismically active to an evidence
 441 based focal point within Egypt's renewable energy strategy, highlighting its
 442 strong potential for sustainable geothermal resource utilization.

443 7. Conclusions

444 The study of 408 micro earthquakes in Abu Dabbab, integrating focal
 445 mechanism analysis with depth dependent stress tensor inversion, reveals a
 446 vertically stratified crust composed of three mechanically distinct layers.
 447 Findings from Abu Dabbab highlight structural complexity and stress
 448 heterogeneity within its seismotectonic and geothermal framework.

449 The shallow crust (0–5 km) is mainly shaped by Red Sea related extension,
 450 with localized compressional effects arising from rigid intrusions and the
 451 complex geometry of reactivated Nadj faults.

452 Intermediate crust (5-10 km): A heterogeneous transition zone where
 453 normal, strike-slip, and reverse faulting coexist, reflecting the interplay of
 454 regional extension and magmatic intrusions.

455 The deep crust (>10 km) is dominated by a relatively uniform oblique
 456 normal extensional regime aligned with the broader regional tectonic field, as
 457 the influence of near surface heterogeneities and magmatic perturbations
 458 diminishes with depth, resulting in a more stable stress environment.

459 Seismicity in Abu Dabbab is shaped by both regional tectonics and local
 460 magmatism. The Red Sea rift imposes NE-SW compression and SE-NW
 461 extension, reactivating Najd faults and influencing stress orientations. A mid
 462 crustal magmatic intrusion further concentrates stress, driving transient stress
 463 rotations, fluid pressure increases, and deformation that sustain swarm activity.

464 In volcanic and geothermal environments, brittle failure can involve the
 465 sudden opening of tensile cracks, producing non double couple radiation
 466 patterns^[18,19]. This dual control mechanism explains the clustered distribution of
 467 seismicity and its depth migration, while also accounting for the anomalously
 468 high heat flow and the coexistence of contrasting stress regimes in close
 469 proximity.

470 In the study area, a multiphasic stress environment complicates seismic
 471 hazard assessment. Faults trending NE-SW or NW-SE may shift between
 472 normal, strike slip, or reverse slip depending on depth and local stress
 473 perturbations. This variability increases uncertainty in seismic source
 474 characterization, making hazard evaluation more complex.

475 In Abu Dabbab, the southern zone at intermediate depths shows
 476 concentrated, heterogeneous seismicity, marking it as an area of elevated
 477 seismic potential that should be prioritized for hazard monitoring. At the same
 478 time, the region demonstrates strong geothermal potential, underscoring its dual
 479 importance for both risk assessment and energy exploration.

480 In Abu Dabbab, persistent seismicity indicates an active mid crustal
 481 magmatic heat source, while widespread normal and oblique faulting reveals a
 482 fractured crust that enables fluid circulation. Together, these conditions create a
 483 favorable setting for geothermal resource development. It also combines
 484 sustained heat, pervasive permeability, and long lived fluid transport, making it
 485 an exceptionally promising site for geothermal exploration and sustainable
 486 energy development.

487 Analysis of 408 microearthquakes in the Abu Dabbab area reveals a
 488 complex stress field characterized by normal, strike-slip, and reverse faulting
 489 across different depth ranges. The coexistence of compressional and extensional
 490 regimes reflects the interaction between NE-SW compression and SE-NW
 491 extension associated with the broader Red Sea rift system, in agreement with

492 previous studies. Recurrent swarms of shallow to intermediate depth seismicity
 493 delineate zones of stress concentration along active faults, providing critical
 494 insights for assessing future earthquake potential and understanding localized
 495 stress perturbations that may trigger seismic events.

496 The predominance of normal and oblique slip faulting in Abu Dabbab
 497 enhances crustal permeability, facilitating fluid circulation and heat transport,
 498 which are key factors for geothermal exploration. This study underscores the
 499 value of high precision earthquake relocation, focal mechanism analysis, and
 500 stress tensor inversion in rift environments. Integrating seismic observations
 501 with geophysical and geochemical data further constrains the roles of magmatic
 502 intrusions and crustal heterogeneity in controlling stress evolution and
 503 seismicity, thereby improving our understanding of the tectono magmatic
 504 processes governing Abu Dabbab's seismic behavior.

505 **8. Implications**

506 The analyses of 407 microearthquakes in the Abu Dabbab region provide
 507 important insights into the seismotectonic and geodynamic behavior of this part
 508 of the Eastern Desert. The identification of multiple faulting mechanisms normal,
 509 strike-slip, and reverse across different depth ranges highlights the complexity of
 510 the local stress field and the coexistence of compressional and extensional
 511 regimes. These findings are consistent with broader Red Sea rift related
 512 tectonics and reflect the interplay between NE-SW compression and SE-NW
 513 extension observed in previous studies.

514 From a seismic hazard perspective, the shallow and intermediate depth
 515 seismicity, particularly the recurrent swarms, points to areas of potential stress
 516 concentration along active faults. Such information is critical for assessing the
 517 likelihood of future earthquake activity and understanding localized stress
 518 perturbations that could trigger seismic events.

519 The observed faulting patterns also have implications for geothermal energy
 520 exploration. The predominance of normal and oblique slip faults, coupled with
 521 evidence of extensional structures, suggests enhanced crustal permeability that
 522 could facilitate fluid circulation and heat transport. Finally, the study
 523 underscores the importance of high resolution earthquake relocation, focal
 524 mechanism determination, and stress tensor inversion in rift environments.
 525 Integrating these seismic observations with geophysical and geochemical
 526 datasets can further constrain the influence of magmatic intrusions and
 527 heterogeneous crustal structures on local stress evolution and seismicity.
 528 Overall, the findings contribute to a more comprehensive understanding of the

529 interplay between tectonic and magmatic processes in shaping Abu Dabbab's
 530 seismic dynamics.

531

532 **Data Availability**

533 The datasets used and/or analysed during the current study are
 534 available from the corresponding author upon reasonable request.

535

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539

540 **Competing Interests**

541 The authors declare no competing interests.

542

543 **Author Contributions**

544 M.A. wrote the manuscript, and all authors reviewed and approved the final
 545 version. M.A. contributed to conceptualization, methodology, software
 546 development, data analysis, manuscript review and editing. H.G. The
 547 corresponding author is responsible for submitting the competing interests
 548 statement on behalf of all authors and contributed to methodology, manuscript
 549 review, data analysis and editing. S.E. contributed to methodology, data
 550 preparation, manuscript review and editing. S.H. is responsible in data
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 552 review. G.E., M.S., M.Y. and M.H. were responsible for data preparation,
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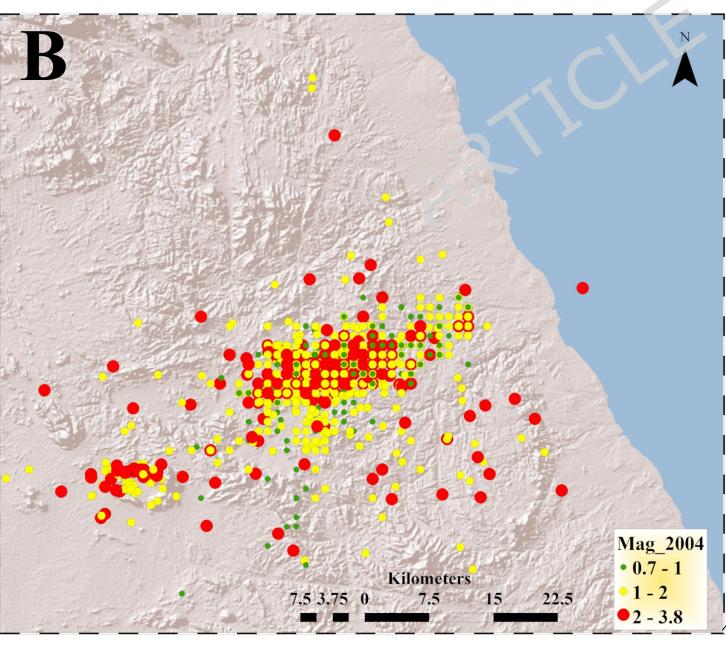
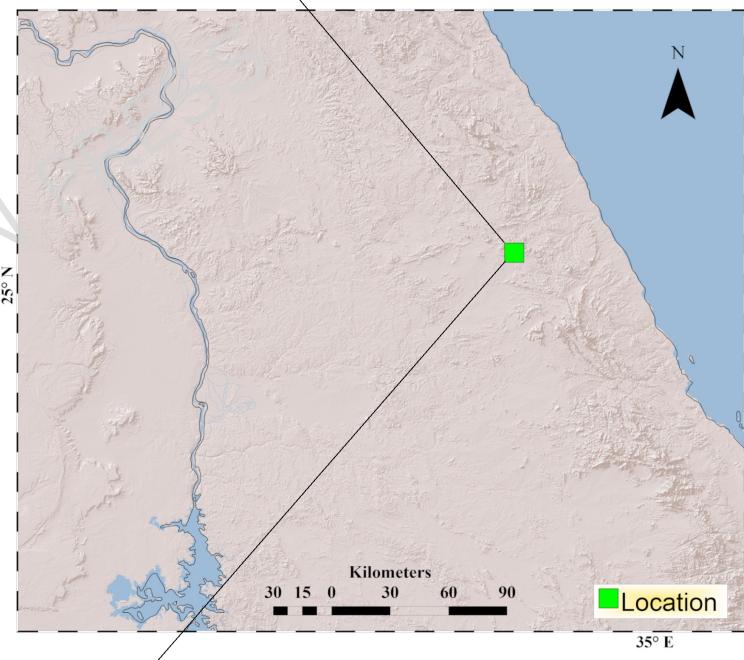
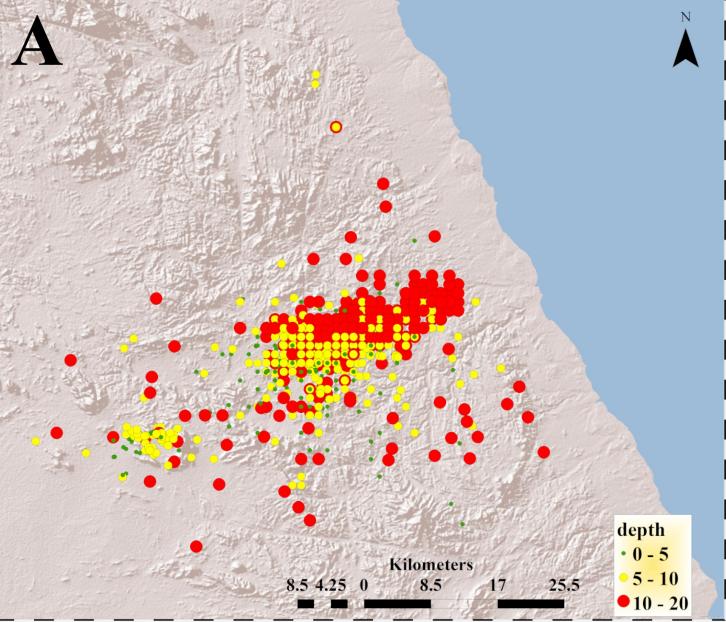
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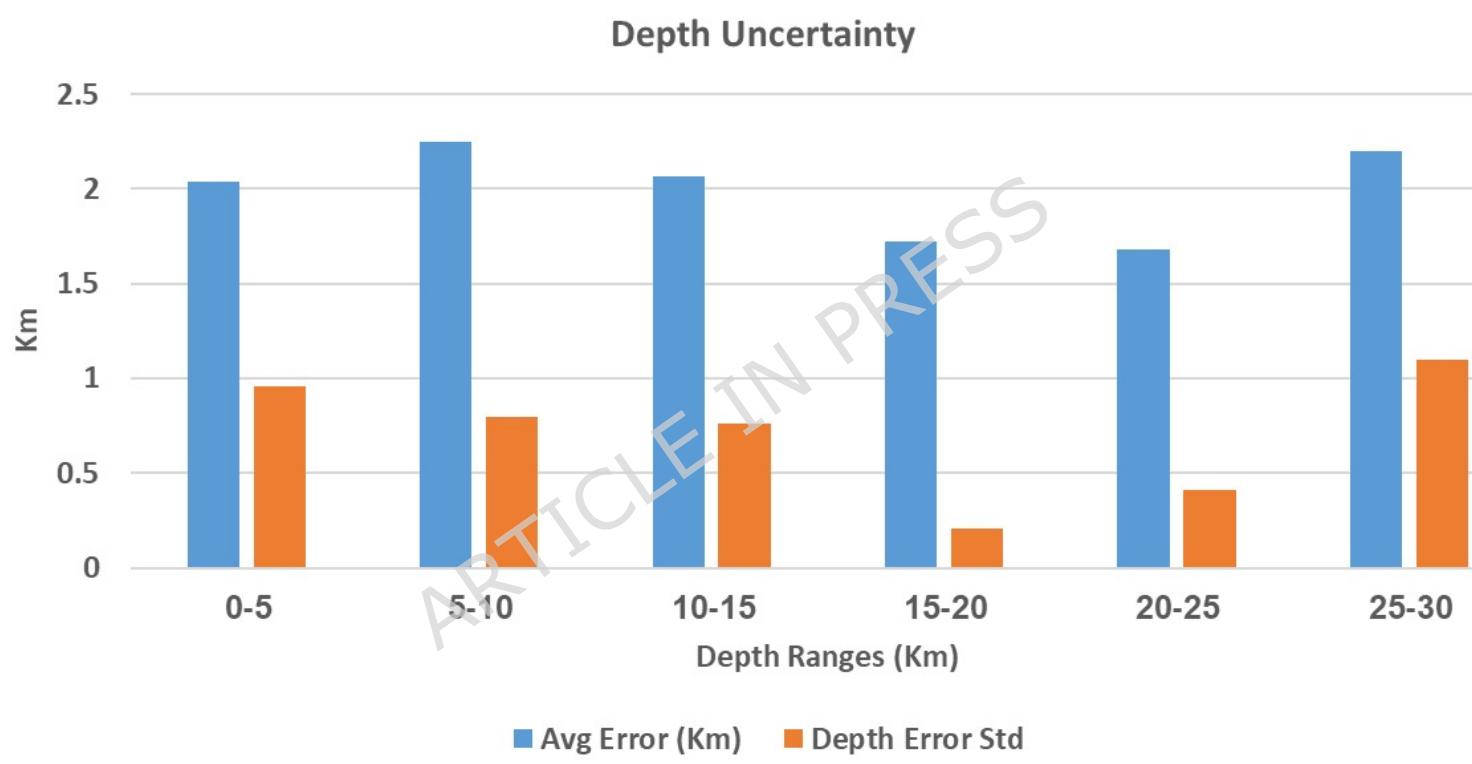
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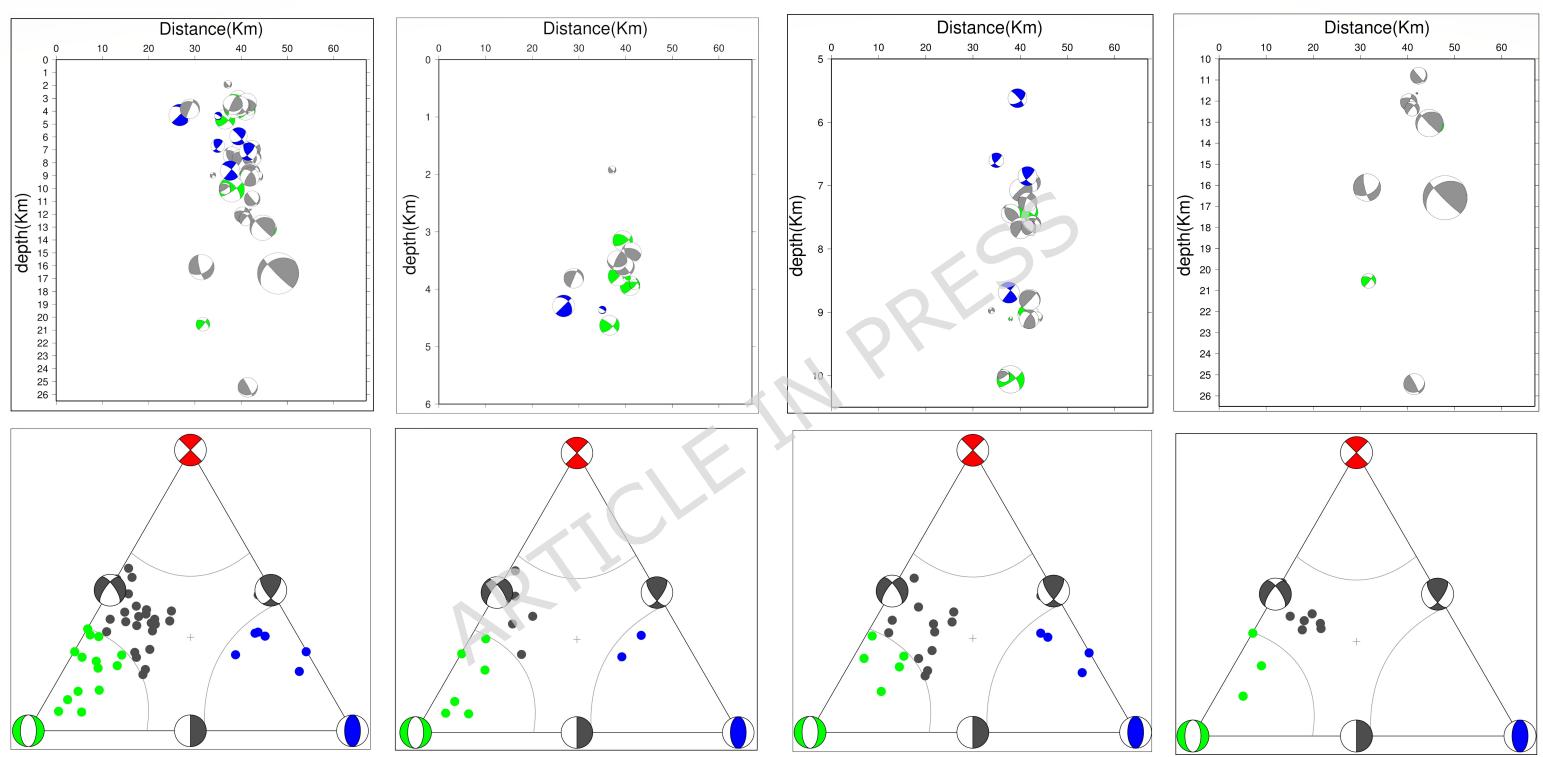
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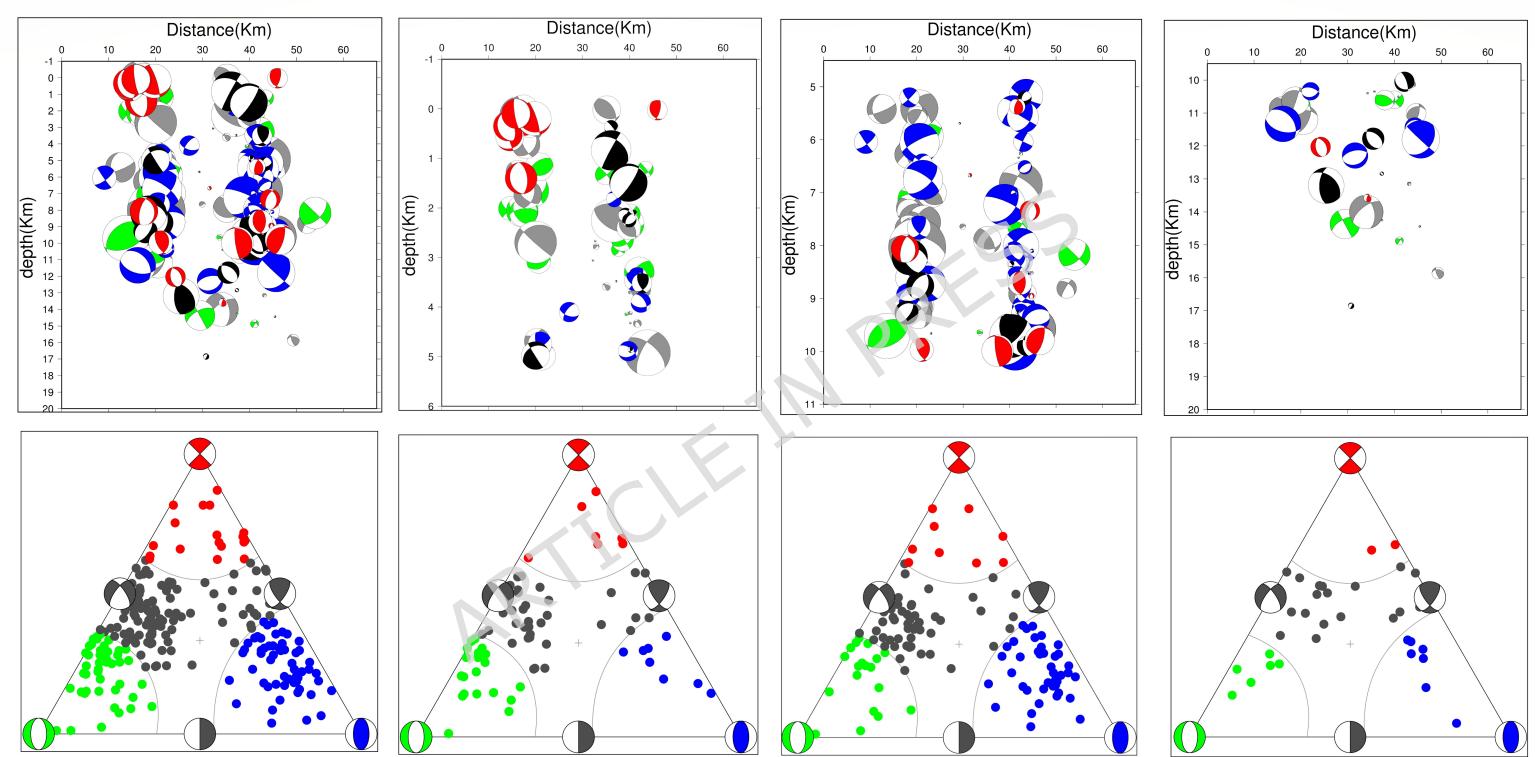
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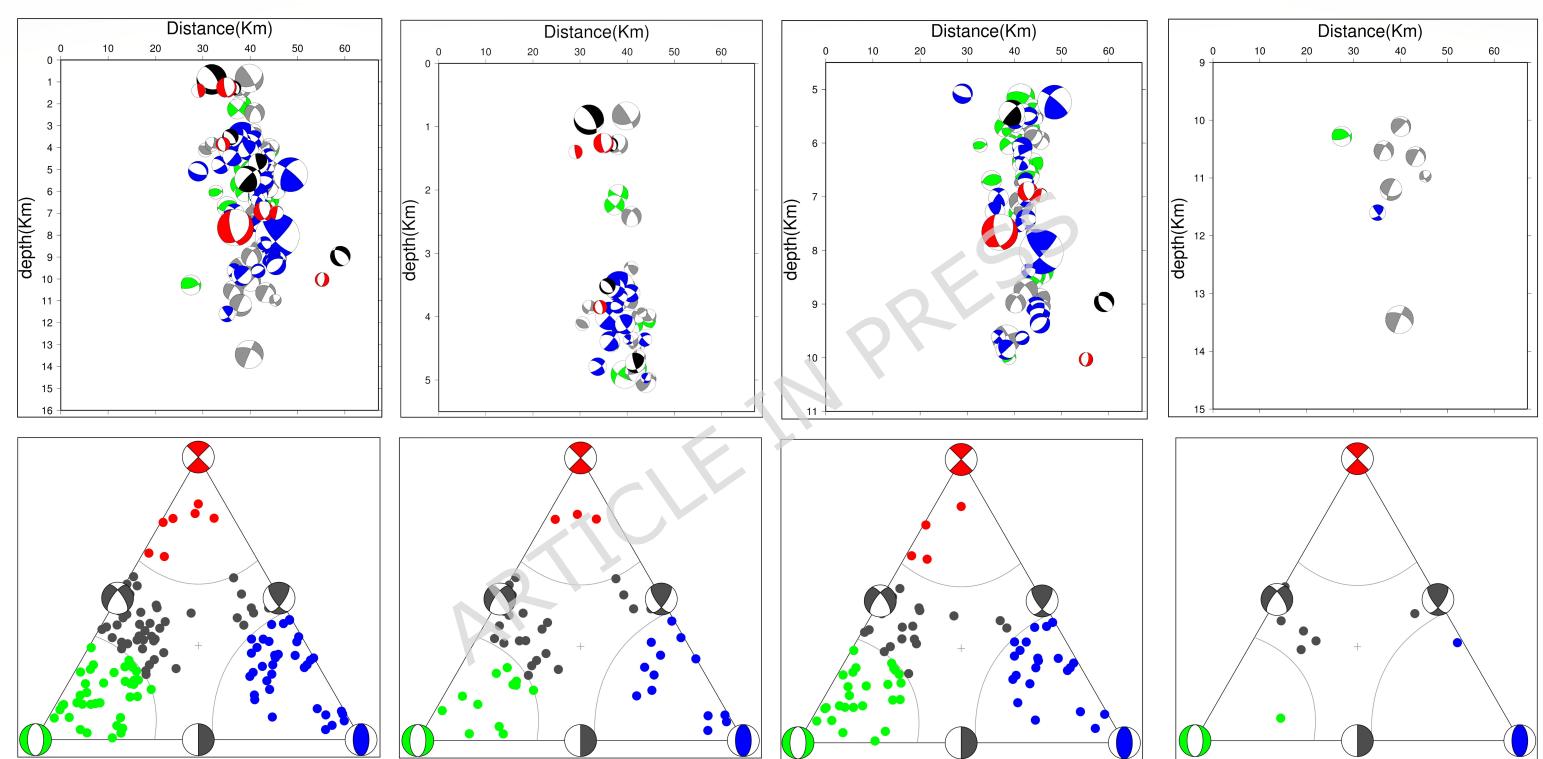


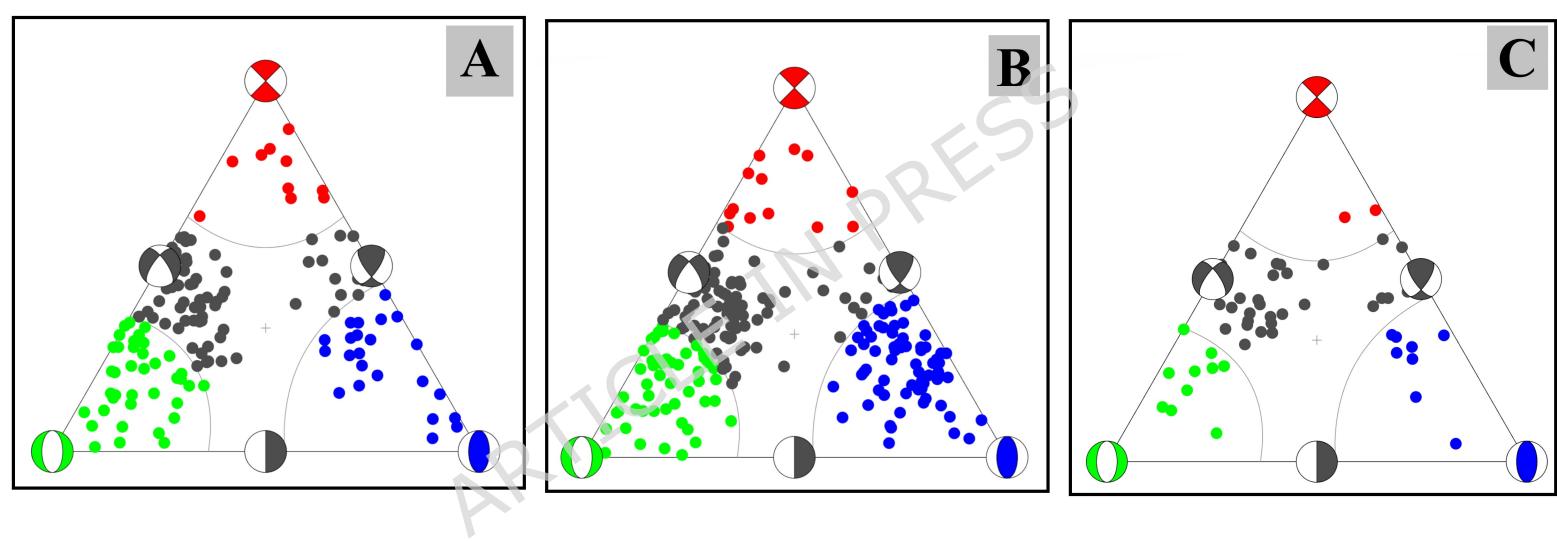


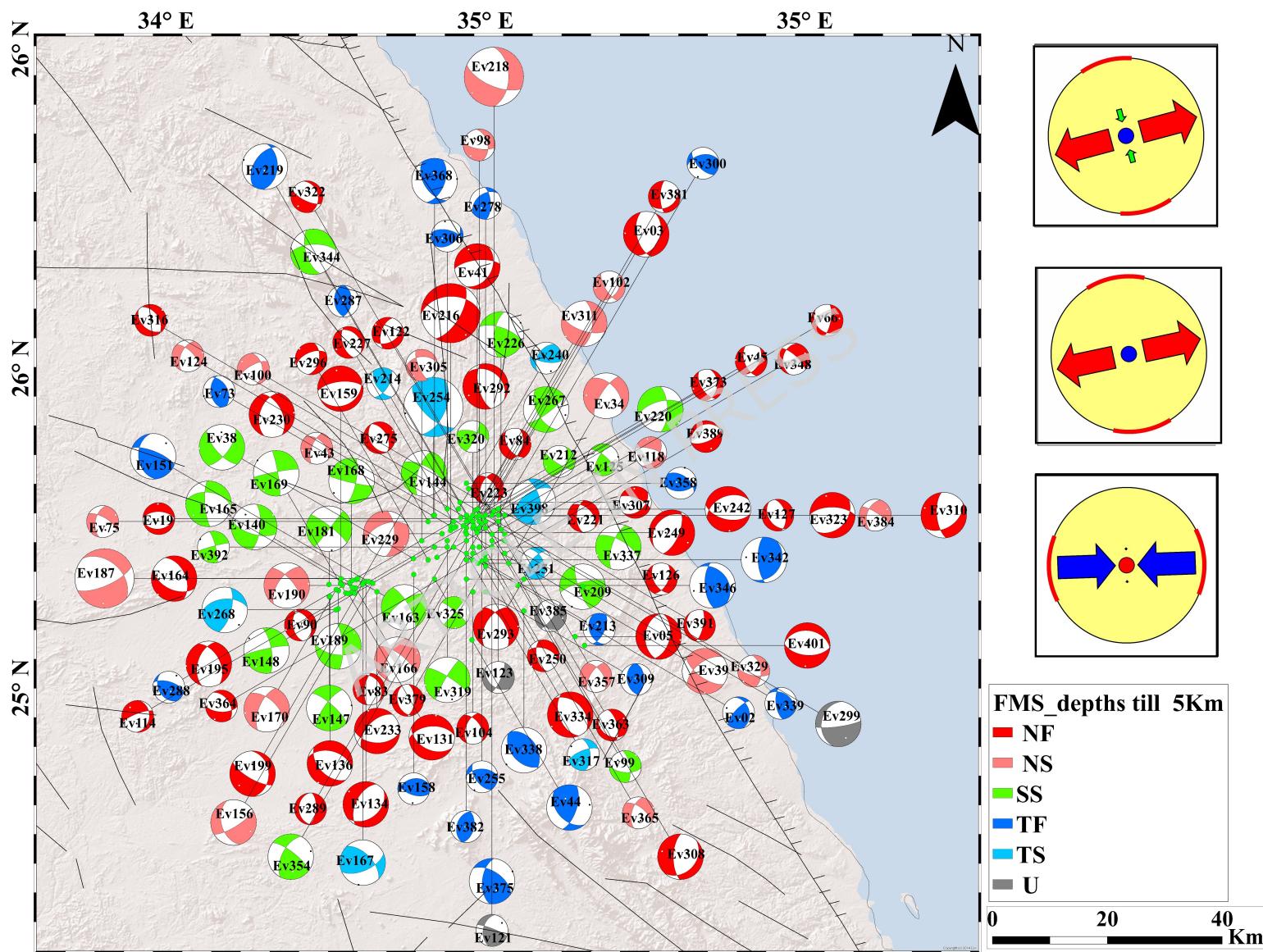


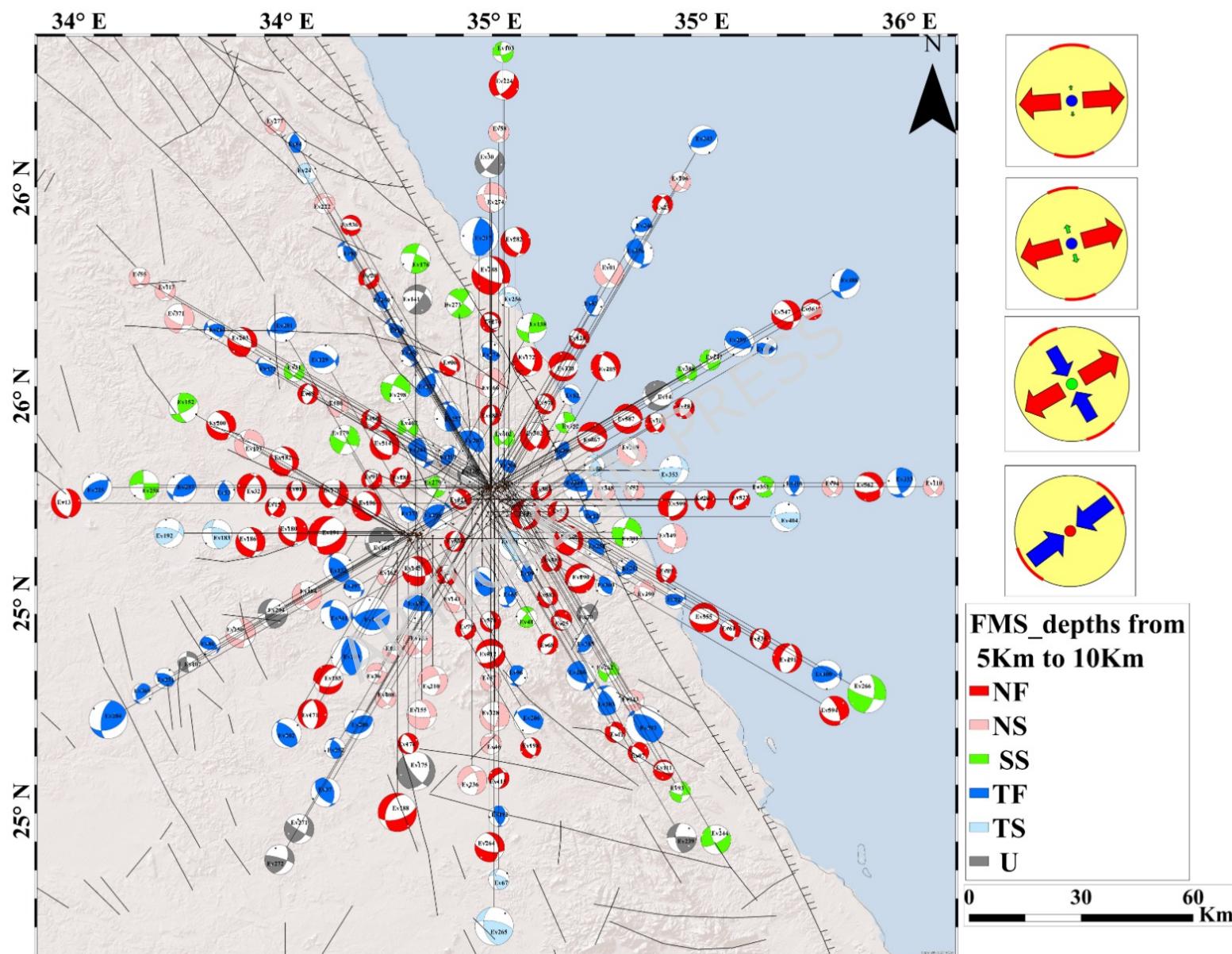


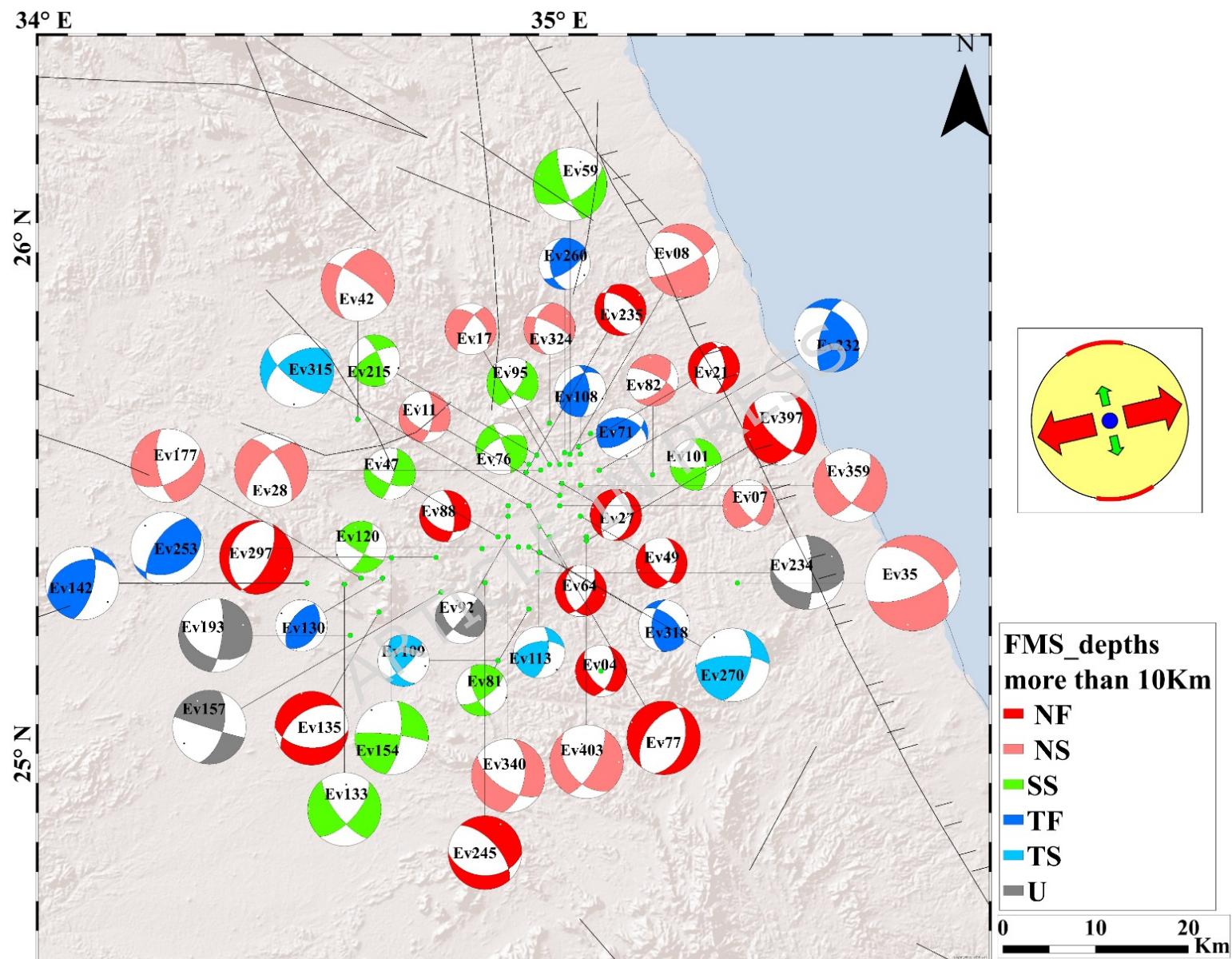
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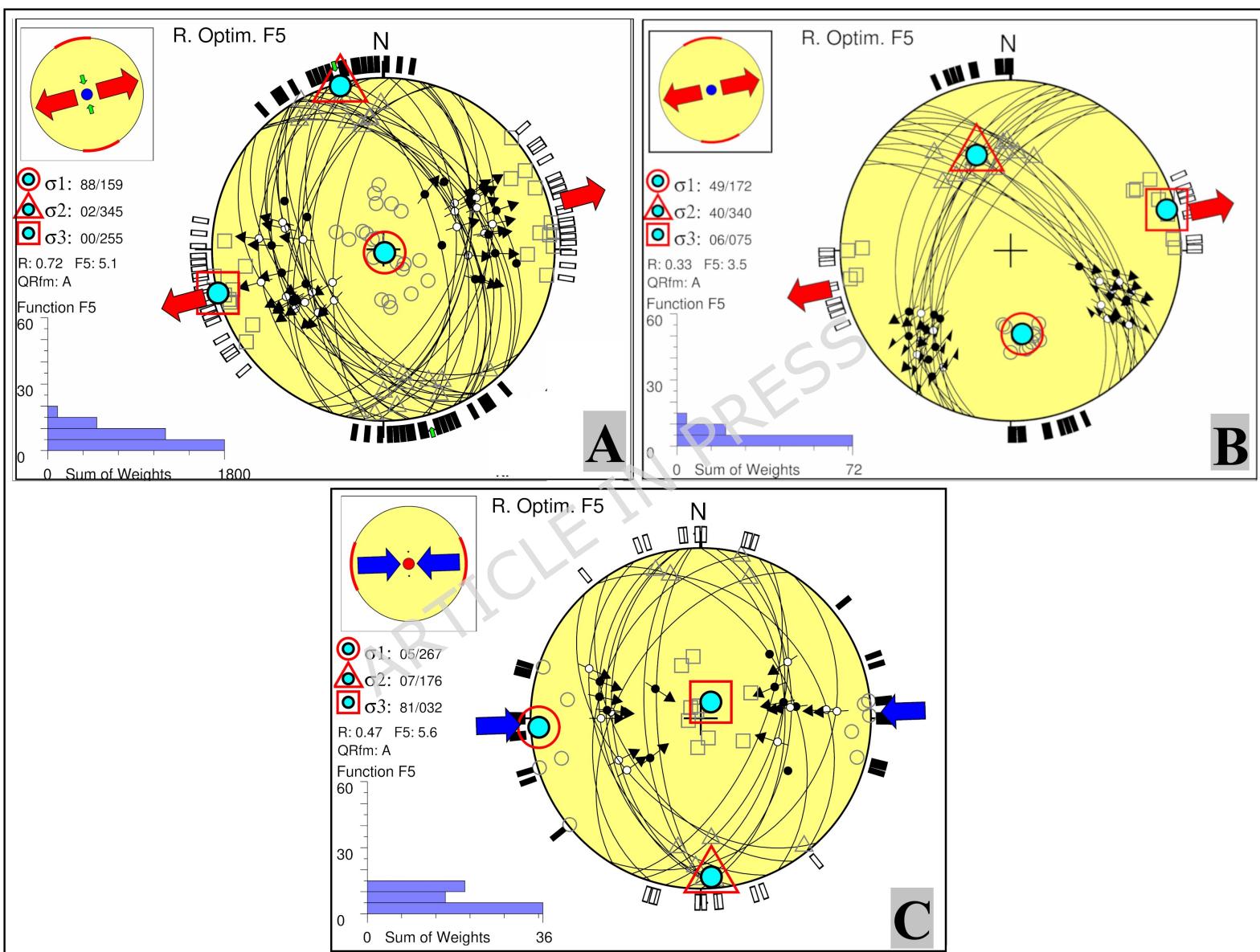


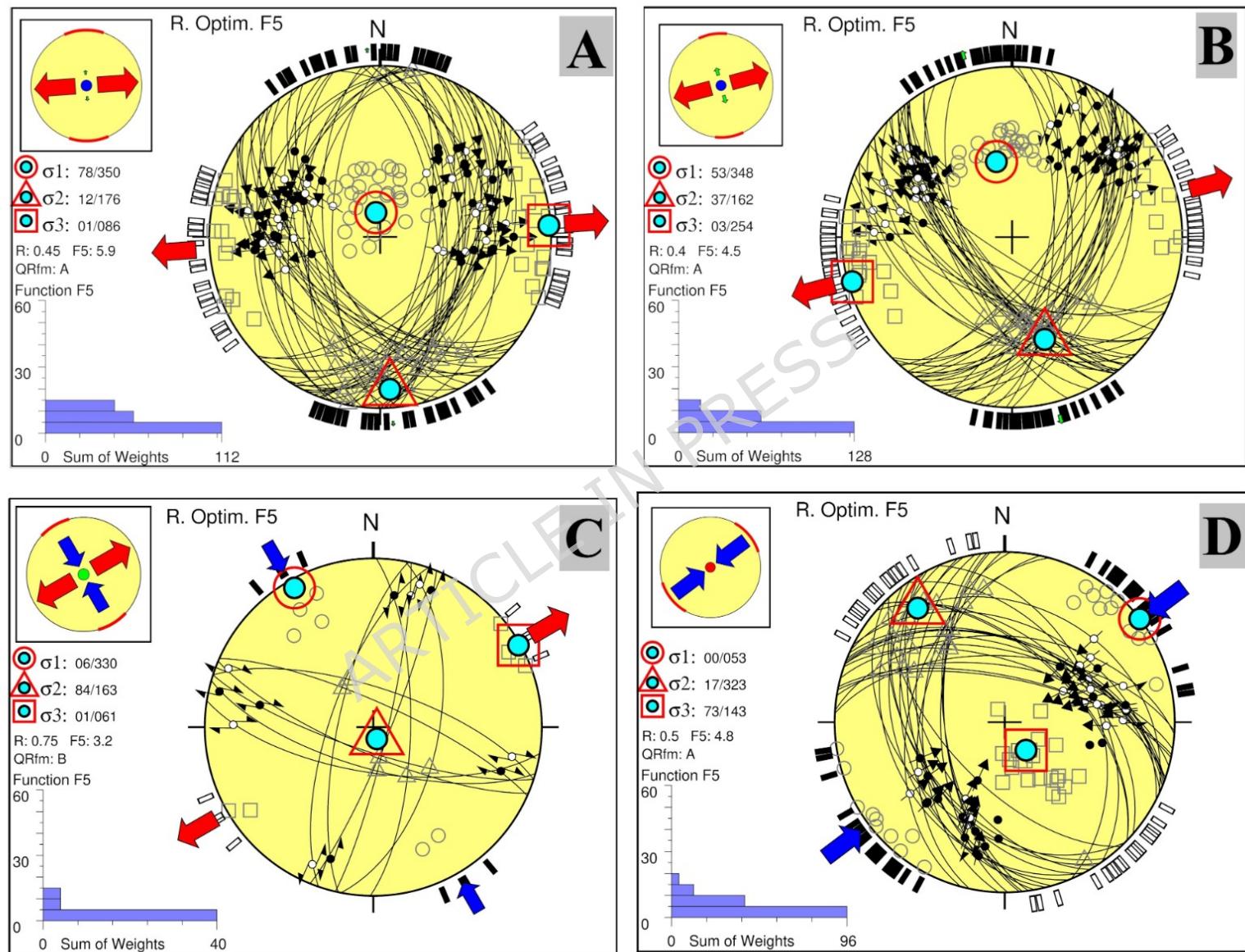


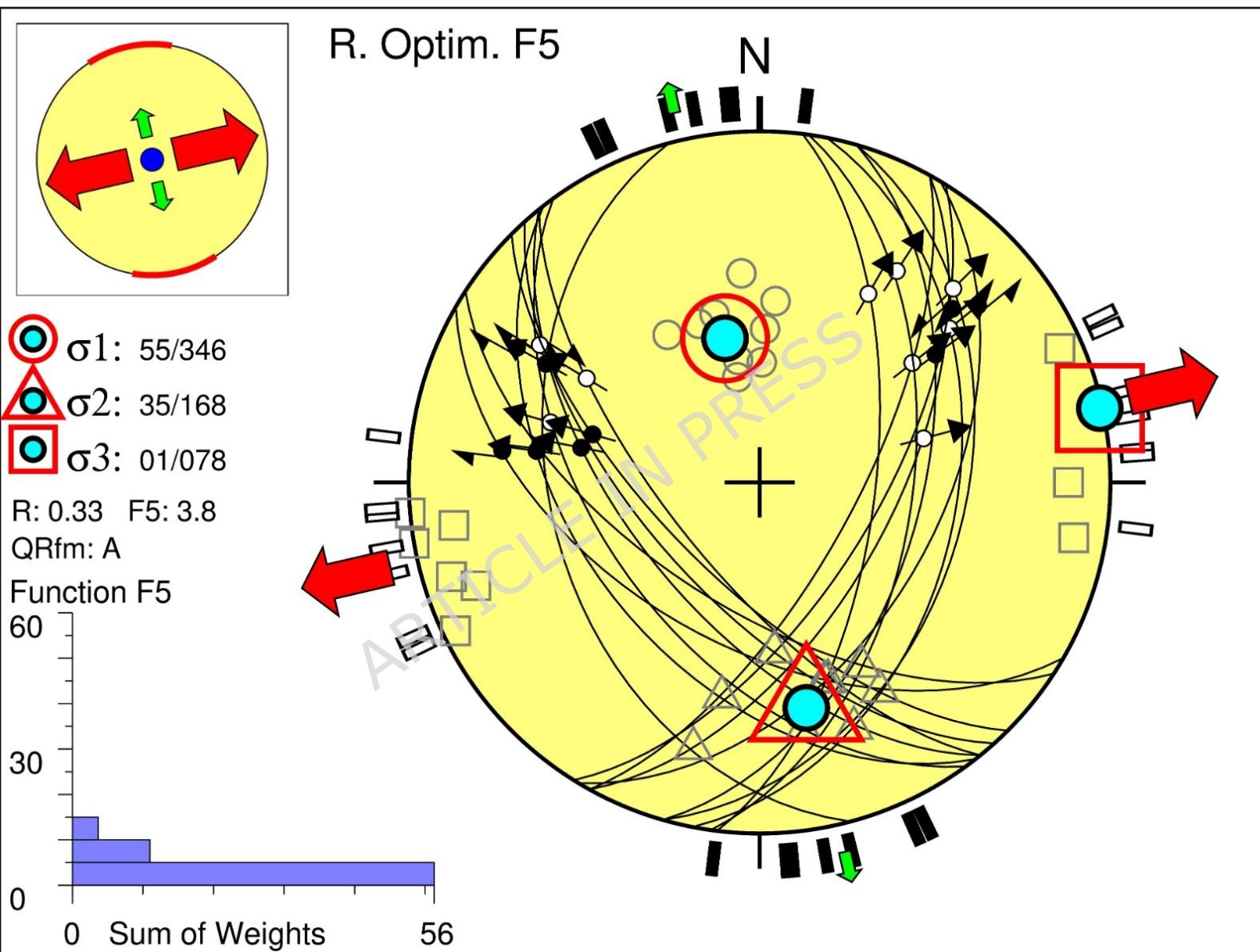












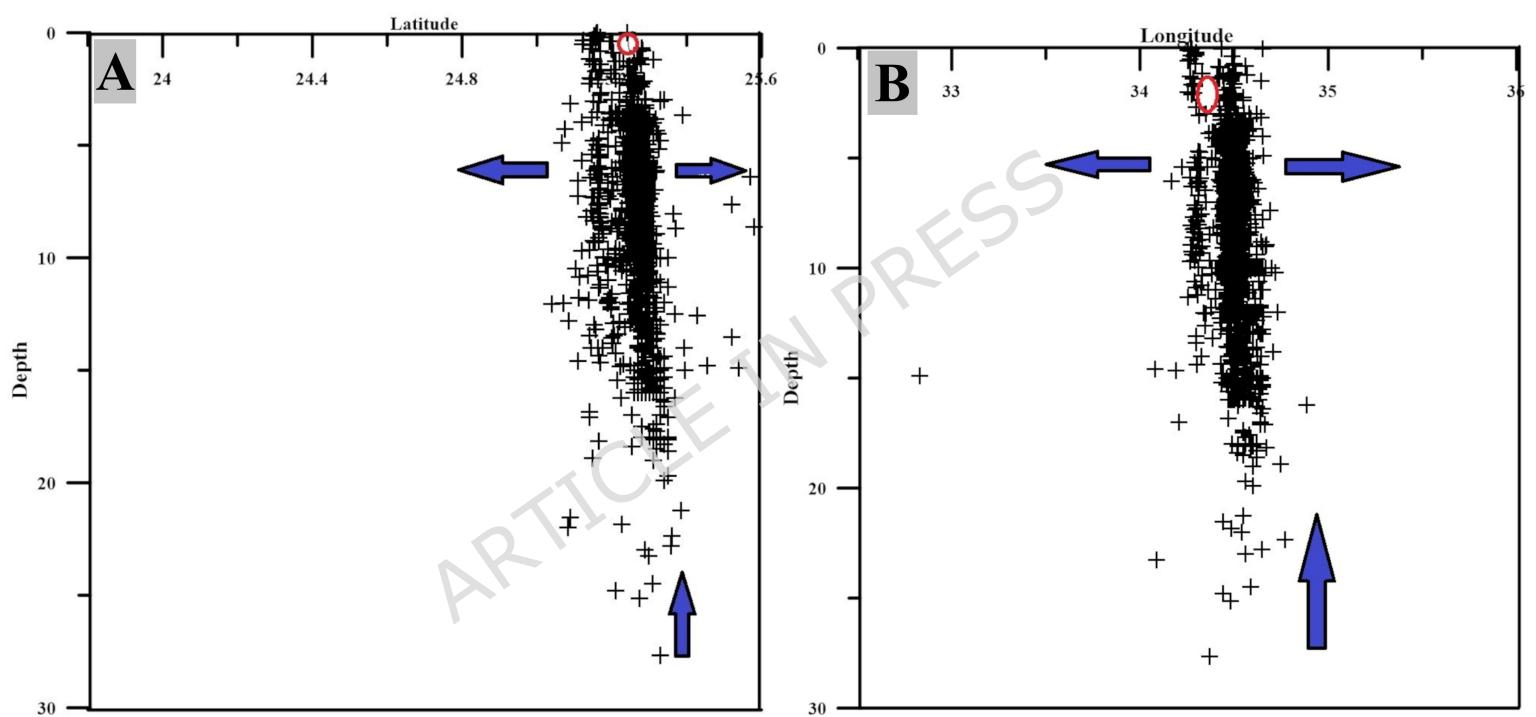


Table 1: Example of hypocentral parameters for selected earthquakes used in constructing focal mechanism solution in Abu Dabbab area

Event_no	year	mon	day	hour	min	sec	long	lat	depth	Ml
Ev01	2004	5	21	0	5	0	34.47	25.27	7.28	1
Ev02	2004	5	21	0	13	0	34.45	25.22	4.36	0.4
Ev03	2004	5	21	2	9	0	34.49	25.26	3.94	1
Ev04	2004	5	21	2	28	0	34.54	25.08	22	0.7
Ev05	2004	5	21	12	33	0	34.64	25.08	3.15	1
Ev06	2004	5	21	16	28	0	34.5	25.24	12.07	0.8
Ev07	2004	5	21	17	15	0	34.51	25.29	12.76	1.3
Ev08	2004	5	21	18	43	0	34.47	25.26	5.62	0.9
Ev09	2004	5	21	20	13	0	34.48	25.27	6.73	0.7
Ev10	2004	5	21	20	54	0	34.5	25.25	12.31	0.7

Table 2: Example of focal mechanism parameters for selected earthquakes in Abu Dabbab area

Event_no	strike_e1	dip1	rake1	strike_e2	dip2	rake2	P_axis		T_axis	
							Az	Pl	Az	Pl
Ev01	143	62	-149	37	63	-32	359	41	90	1
Ev02	344	33	36	222	72	118	292	22	167	55
Ev03	23	40	-61	167	55	-113	25	70	273	8
Ev04	16	52	-51	144	52	-128	350	61	260	0
Ev05	13	49	-84	183	41	-97	330	84	99	4
Ev06	31	48	-29	141	69	-134	5	47	262	13
Ev07	63	75	-133	318	45	-21	292	43	176	18
Ev08	331	49	61	191	49	119	261	0	171	68
Ev09	22	46	-61	163	51	-116	10	69	272	3
Ev10	7	64	-44	120	52	-146	328	49	66	7

Table 3: The results of stress tensor inversion in Abu Dabbab area

At shallow depth (0-5 Km)													
(group_A)													
σ_1		σ_2		σ_3		R	α	R'	F5	quality	SH_{max}	Sh_{min}	Stress regime
Az	Pl	Az	Pl	Az	Pl								

15 9	8 8	34 5	0 2	25 5	0 0	0.7 2	10. 3	0.7 2	5. 1	A	165	N75° E	NF
(group_B)													
σ₁	σ₂		σ₃		R	α	R`	F5	qualit y	SH_{ma} x	Sh_{min}	Stress regim e	
Az	Pl	Az	Pl	Az	Pl								
17 2	4 9	34 0	4 0	75	0 6	0.3 3	8.4	1.0	3. 5	A	168	N78° E	NS
(group_C)													
σ₁	σ₂		σ₃		R	α	R`	F5	qualit y	SH_{ma} x	Sh_{min}	Stress regim e	
Az	Pl	Az	Pl	Az	Pl								
26 7	0 5	17 6	0 7	32	8 1	0.4 7	10. 1	2.5	5. 6	A	88	178°	TF
Intermediate depths (5-10 Km)													
(group_A)													
σ₁	σ₂		σ₃		R	α	R`	F5	qualit y	SH_{ma} x	Sh_{min}	Stress regim e	
Az	Pl	Az	Pl	Az	Pl								
35 0	7 8	17 6	1 2	86	0 1	0.4 5	11. 4	0.4 5	5. 9	A	176	86°	NF
(group_B)													
σ₁	σ₂		σ₃		R	α	R`	F5	qualit y	SH_{ma} x	Sh_{min}	Stress regim e	
Az	Pl	Az	Pl	Az	Pl								
34 8	5 3	16 2	3 7	25 4	0 3	0.4	9.5	1	4. 5	A	165	N75° E	NS
(group_C)													
σ₁	σ₂		σ₃		R	α	R`	F5	qualit y	SH_{ma} x	Sh_{min}	Stress regim e	
Az	Pl	Az	Pl	Az	Pl								
33 0	0 6	16 3	8 4	61	1	0.7 5	7.9	1.5	3. 2	B	150	60°	SS
(group_D)													
σ₁	σ₂		σ₃		R	α	R`	F5	qualit y	SH_{ma} x	Sh_{min}	Stress regim e	
Az	Pl	Az	Pl	Az	Pl								
53 0	0 3	32 7	1 3	14 3	7 3	0.5	10. 6	2.5	4. 8	A	53	143°	TF
At deeper depths (more than 10 Km)													
σ₁	σ₂		σ₃		R	α	R`	F5	qualit y	SH_{ma} x	Sh_{min}	Stress regim e	
Az	Pl	Az	Pl	Az	Pl								
34 6	5 5	16 8	3 5	78	0 1	0.3 3	9	01	3. 8	A	167	N77° E	NS